

**The Direct Sampling
technique**

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Simulation of rainfall time-series from different climatic regions using the Direct Sampling technique

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Abstract

The Direct Sampling technique, belonging to the family of multiple-point statistics, is proposed as a non-parametric alternative to the classical autoregressive and Markov-chain based models for daily rainfall time-series simulation. The algorithm makes use of the patterns contained inside the training image (the past rainfall record) to reproduce the complexity of the signal without inferring its prior statistical model: the time-series is simulated by sampling the training dataset where a sufficiently similar neighborhood exists. The advantage of this approach is the capability of simulating complex statistical relations by respecting the similarity of the patterns at different scales. The technique is applied to daily rainfall records from different climate settings, using a standard setup and without performing any optimization of the parameters. The results show that the overall statistics as well as the dry/wet spells patterns are simulated accurately. Also the extremes at the higher temporal scale are reproduced exhaustively, reducing the well known problem of over-dispersion.

1 Introduction

The stochastic generation of rainfall time-series is a key topic for hydrological and climate science applications: the challenge is to simulate a synthetic signal honoring the high-order statistics observed in the historical record, respecting the seasonality and persistence from the daily to the higher temporal scales. Among the different proposed techniques, exhaustively reviewed by Sharma and Mehrotra (2010), the most commonly used approach to the problem, adopted since the '60, is the Markov-chain (MC) simulation: in its classical form, it is a linear model which cannot simulate the variability and persistence at different scales. Some recently adopted solutions to overtake this limit consist of introducing exogenous climatic variables and large-scale circulation indexes (Hay et al., 1991; Bardossy and Plate, 1992; Katz and Parlange, 1993; Woolhiser et al., 1993; Hughes and Guttorp, 1994; Hughes et al., 1999; Wallis and Griffiths, 1997;

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Wilby, 1998; Kiely et al., 1998), lower-frequency daily rainfall covariates (Wilks, 1989; Briggs and Wilks, 1996; Jones and Thornton, 1997; Katz and Zheng, 1999) or an index based on the short-term daily historical or previously generated record (Harrold et al., 2003a, b; Mehrotra and Sharma, 2007a, b) as conditional variables for the estimation of the MC parameters. By doing this, non-linearity is introduced in the prior model, the MC parameters changing in time as a function of some specific low-frequency fluctuations. An alternative proposed method is model nesting (Wang and Nathan, 2002; Srikanthan, 2004, 2005; Srikanthan and Pegram, 2009), that implies the correction of the generated daily rainfall using a multiplicative factor to compensate the bias in the higher-scale statistics. These techniques generally allow a better reproduction of the statistics up to the annual scale, but they imply the estimation of a more complex prior model and cannot completely catch a complex dependence structure.

In this paper, we propose the use of some lower-frequency covariates of daily rainfall in a completely unusual framework: the Direct Sampling (DS) technique (Mariethoz et al., 2010), which belongs to multiple-point statistics (MPS). Introduced by Guardiano and Srivastava (1993) and widely developed during the last decade (Strebelle, 2002; Allard et al., 2006; Zhang et al., 2006; Arpat and Caers, 2007; Honarkhah and Caers, 2010; Straubhaar et al., 2011; Tahmasebi et al., 2012), MPS is a family of geostatistical techniques widely used in spatial data simulations and particularly suited to pattern reproduction. MPS algorithms use a training image, i.e. a dataset to evaluate the probability distribution (pdf) of the variable simulated at each point (in time or space), conditionally to the values present in its neighborhood. In the particular case of the Direct Sampling, the concept of training image is taken to the limit by avoiding the computation of the conditional pdf and making a random sampling of the historical dataset where a pattern similar to the conditional data is found. If the training dataset is representative enough, these techniques can easily reproduce high-order statistics of complex natural processes at different scales. MPS has already been successfully applied to the simulation of spatial rainfall occurrence patterns (Wojcik et al., 2009). In this paper, we test the Direct Sampling on the simulation of daily rainfall time-series. The

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aim is to reproduce the complexity of the rainfall signal up to the decennial scale, simulating the occurrence and the amount at the same time with the aid of a multivariate dataset. Similar algorithms performing a multivariate simulation had been previously developed by Young (1994) and Rajagopalan and Lall (1999) using a bootstrap-based approach. As discussed in details in Sect. 2.3, the advantage of the Direct Sampling with respect to the mentioned techniques is the possibility to have a variable high-order time-dependence, without incurring excessive computation since the estimation of the n dimensional conditional pdf is not needed. Moreover, we propose a standard setup for rainfall simulation: an ensemble of auxiliary variables and fixed values for the main parameters required by the Direct Sampling algorithm, suitable for the simulation of any stationary rainfall time-series, without the need of calibration. The technique is tested on three time-series from different climatic regions of Australia. The paper is organized as follows: in Sect. 2 the DS technique, the dataset used and the method of evaluation are described. The statistical analysis of the simulated time-series is presented and discussed in Sect. 4. Section 5 is dedicated to the conclusions.

2 Methodology

In this section we recall the basics of multiple-point statistics and we focus on the Direct Sampling algorithm. The dataset used is then presented as well as the methods of evaluation.

2.1 Background on multiple-point statistics

Before entering in the details of the DS algorithm, let us introduce some common elements of MPS. The whole information used by MPS to simulate a certain process is based on the *Training Image* (TI) or *training dataset*: the dataset constituted of one or more variables used to infer the statistical relations and occurrence probability of any datum in the simulation. The TI may be constituted of a conceptual model instead of

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real data, but in the case of the rainfall time-series it is more likely to be a historical record of rainfall measurements. The *Simulation Grid* (SG) is a totally or partially un-informed N-dimensional array in which the algorithm generates values to obtain the actual output of the simulation. It usually has the same dimensionality as the training image. In the case of a rainfall time-series simulation, the SG is a time-referenced one-dimensional vector of random variables (referred to as one variable for the sake of simplicity), each of which represents the rainfall amount for a certain time-step. The *conditioning data* (CD) are a group of given data (e.g. rainfall measurements) situated in the SG. Being already informed, no simulation occurs at those time-steps. The presence of CD affects, in their neighborhood, the conditional law used for the simulation and limits the range of possible patterns. MPS, as well some MC-based algorithm for rainfall simulation (see Sect. 1), may include the use of *auxiliary variables* to condition the simulation of the target variable. An auxiliary variable is normally given as CD but, in the case of the Direct Sampling, it can also be co-simulated with the target, without being necessarily informed. For rainfall time-series, it could be for example: covariates of the original or previously simulated data (e.g. the number of wet days in a past period), a correlated variable for which the record is known, a theoretical variable that imposes a periodicity or a trend (e.g. a sinusoid function describing the annual seasonality over the data). Finally, the *search neighborhood* is a moving window, i.e. the portion of time-series located in the past and future neighborhood of each simulated value, used to retrieve the *data event*, i.e. the group of time-referenced values used to condition the simulation.

2.2 The Direct Sampling algorithm

Classical MPS implementations create a catalog of the possible neighbors patterns to evaluate the conditional probability of occurrence for each event with respect to the considered neighborhood. This may imply a significant amount of memory and always limits the application to categorical variables. On the contrary, the Direct Sampling generates each value by sampling the data from the TI where a sufficiently similar

neighborhood exists. The DS implementation used in this paper is called *DeeSse* software (Straubhaar, 2011), the following is the main workflow of the algorithm. Let us denote $\mathbf{x} = [x_1, \dots, x_n]$ the time vector representing the SG, $\mathbf{y} = [y_1, \dots, y_m]$ the one representing the TI and $Z(\cdot)$ the target variable, object of the simulation, defined at each element of \mathbf{x} and \mathbf{y} . Before the simulation begins, all continuous variables are normalized using the transformation $Z \mapsto Z \cdot (\max(Z) - \min(Z))^{-1}$ in order to have distances (see step 3) in the range $[0, 1]$. During the simulation, all the time-steps of the SG are visited in a random order. At each uninformed x_t , the following steps are executed:

1. The data event $\mathbf{d}(x_t) = (Z(x_{t+h_1}), \dots, Z(x_{t+h_n}))$ is retrieved from the SG according to a fixed neighborhood of radius R centered on x_t . It consists of at most N informed time-steps, closest to x_t and defines a set of lags $H = \{h_1, \dots, h_n\}$, with $|h_j| \leq R$ and $n \leq N$. The size of $\mathbf{d}(x_t)$ is therefore limited by the user-defined parameter N and the available informed time-steps inside the search neighborhood.
2. A random time-step y_j in \mathbf{y} is visited and the corresponding data event $\mathbf{d}(y_j)$, defined according to the same H , is retrieved to be compared with $\mathbf{d}(x_t)$.
3. A distance $D(\mathbf{d}(x_t), \mathbf{d}(y_j))$, i.e. a measure of dissimilarity between the two data events, is calculated. For categorical variables (e.g. the dry/wet rainfall sequence), it is given by the formula:

$$D(\mathbf{d}(x_t), \mathbf{d}(y_j)) = \frac{1}{n} \sum_{j=1}^n a_j, \quad a_j = \begin{cases} 1 & \text{if } Z(x_j) \neq Z(y_j) \\ 0 & \text{if } Z(x_j) = Z(y_j) \end{cases} \quad (1)$$

while for continuous variables the following one is used:

$$D(\mathbf{d}(x_t), \mathbf{d}(y_j)) = \frac{1}{n} \sum_{j=1}^n |Z(x_t) - Z(y_j)| \quad (2)$$

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where n is the number of elements of the data event. The elements of $\mathbf{d}(x_t)$, independently from their position, play an equivalent role in conditioning the simulation of $Z(x_t)$. Note that, using the above distance formulas, normalization is not needed for categorical variables, while for the continuous ones it ensures distances in $[0, 1]$.

4. If $D(\mathbf{d}(x_t), \mathbf{d}(y_i))$ is below a fixed threshold T , i.e. the two data events are sufficiently similar, the iteration stops and the datum $Z(y_i)$ is assigned to $Z(x_t)$. Otherwise, the process is repeated from point 2 until a suitable candidate $\mathbf{d}(y_i)$ is found or the prescribed TI fraction limit F is scanned.
5. If a TI fraction F has been scanned and the distance $D(\mathbf{d}(x_t), \mathbf{d}(y_i))$ is above T for each visited y_i , the datum $Z(y_i^*)$ minimizing this distance is assigned to $Z(x_t)$.

This procedure is repeated for the simulation at each x_t until the entire SG is informed. Figure 1 illustrates the iterative simulation using the Direct Sampling and stresses some of its peculiarities. First, simulating $Z(x_t)$ in a random order allows \mathbf{x} to be progressively populated at non-consecutive time-steps. Therefore, the simulation at each x_t can be conditioned on both past and future, as opposed to the classical Markov-chain techniques, that use a linear simulation path starting from the beginning of the series, allowing conditioning on past only.

In the early iterations, the closest informed time-steps used to condition the simulation are located far from x_t and its number is limited by the search window, i.e. conditioning is mainly based on large past and future time-lags. On the contrary, the final iterations dispose of a more populated SG, conditioning is thus done on small time-lags since only the closest N values are considered. This variable time-lag principle may not respect the autocorrelation on a specific time-lag rigorously, but it should reproduce a more complex statistical relationship, which cannot be explored exhaustively using a fixed-dependence model.

The DS can simulate multiple variables together similarly to the univariate case, but using a multivariate dataset as TI. In this case, we have a vector of variables $Z(x_t)$

defined at each time-step. Point 3 of the algorithm is therefore repeated for each variable using an independent data event and acceptance threshold. It has to be remarked that the entire data vector $Z(x_t)$ is simulated in one iteration, reproducing exactly the same combination of values found for all the variables at the sampled time-step, excluded the already informed ones (conditioning data). This feature, although reducing the variability in the simulation, has been adopted to reproduce the correlation between variables accurately.

2.3 Comparison with existing resampling techniques

The resampling principle is at the base of some already proposed techniques for rainfall and hydrologic time-series simulation. There exist two principal families of resampling techniques: the block bootstrap (Vogel and Shallcross, 1996; Srinivas and Srinivasan, 2005; Ndiritu, 2011), which implies the resampling with replacement of entire pieces of time-series with the aim of preserving the statistical dependence at a scale smaller than the blocks size, and the k -nearest neighbor bootstrap (k -NN), based on single value resampling using a pattern similarity rule. This latter family of techniques, introduced by Efron (1979) and inspired to the jackknife variance estimation, has seen several developments in hydrology (Young, 1994; Lall and Sharma, 1996; Lall et al., 1996; Rajagopalan and Lall, 1999; Buishand and Brandsma, 2001; Wojcik and Buishand, 2003; Clark et al., 2004). Having different points in common with the Direct Sampling, its general framework is briefly presented in the following. Each datum inside the historical record is characterized by a vector \mathbf{d}_t of predictor variables, analogous to the data event for the DS. For example, to generate $Z(x_t)$ one could use $\mathbf{d}_t = [Z(x_{t-1}), Z(x_{t-2}), U(x_t), U(x_{t-1})]$, meaning that the simulation is conditioned to the 2 previous time-steps of Z and the present and previous time-steps of U , a correlated variable. In the predictor variables space ID, the historical data as well as $Z(x_t)$, which still has to be generated, are represented as points whose coordinates are defined by \mathbf{d}_t . Consequently, proximity in ID corresponds to similarity of the conditioning patterns. $Z(x_t)$ is simulated by sampling an empirical pdf constructed on the k points closest to

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$Z(x_t)$; the closer the point is, the higher is the probability to sample the corresponding historical datum. Proposed variations of the algorithm include transformations of the predictor variables space, the application of kernel smoothing to the k -NN pdf to increase the variability beyond the historical values, and different methods to estimate the parameters of the model, e.g. k and the kernel bandwidth.

Going back to the Direct Sampling, the similarities with the k -NN bootstrap are: (i) they both make a resampling of the historical record conditioned by an ensemble of auxiliary/predictor variables; (ii) they both compute a distance as a measure of dissimilarity between the simulating time-step and the candidates considered for resampling. Nevertheless, there are several points of divergence in the rationale of the techniques: (i) in the k -NN bootstrap, the distance is used to evaluate the resampling probability, while in the DS it is used to evaluate the resampling possibility. This means that, using the k -NN resampling, the conditional pdf is a function of the distance, while in the DS the distance is only used to define its support. In fact, using the DS, the space ID is not restricted to the k nearest neighbors but it is bounded by the distance thresholds: outside the boundary, the resampling probability is zero, while inside, it follows the occurrence of the data in the scanned TI fraction, without being a function of the pattern resemblance. Only in case of no candidate found, the closest neighbor outside the bounded portion of ID is chosen for resampling. The latter can be considered as an exceptional condition which usually does not lead to a good simulation and seldom occurs using an appropriate setup and training dataset. (ii) Using the DS, the conditional pdf remains implicit, its computation is not needed: the historical record is randomly visited instead and the first datum presenting a distance below the threshold is sampled. This is an advantage since it avoids the problem of the high-dimensional conditional pdf estimation which limits the degree of conditioning in bootstrap techniques (Sharma and Mehrotra, 2010). (iii) The k -NN technique considers a fixed time-dependence, while it varies during the simulation in the case of the DS. (iv) Finally, the simulation path (in the SG) is always linear in the k -NN technique, while it is random using the DS, allowing conditioning on future time-steps of the target variable.

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3 Application

The dataset chosen for this study is composed of three daily rainfall time-series from different climatic regions of Australia: Alice Springs (hot desert), with a very dry rainfall regime and long droughts, Sydney (temperate), with a far wetter climate due to its proximity to the ocean, and Darwin (tropical savannah), showing and extreme variability between the dry and wet seasons.

Table 1 delineates the dataset used: the chosen stations present a considerable record of about 70 years for Darwin and Alice Springs and 150 years for Sydney. Any gaps or trends have been explicitly kept to test the behavior of the algorithm with incomplete or non-stationary datasets. The Direct Sampling treats gaps in the time-series in a simple way: each data event found in the TI is rejected if it contains any missing data. This allows incomplete training images to be dealt with in a safe way, but, as one could expect, a large quantity of missing data, especially if sparsely distributed, may lead to a poor simulation. About dataset reconstruction using the direct sampling see Mariethoz and Renard (2010).

Since rainfall is a complex signal exhibiting not only multi-scale time dependence but also intermittence, the classical approach is to split the daily time-series generation in two steps: the occurrence model, where the dry/wet daily sequence is generated using a Markov-chain, and the amount model, where the rainfall amount is simulated on wet days using an estimation of the conditional pdf (e.g., Coe and Stern, 1982). The simulation framework proposed here is radically different: we use the Direct Sampling to generate the complete time-series in one step, simulating multiple variables together. In particular, the TI used is composed of the past daily rainfall record (*) and the following auxiliary variables (Table 2): (1) the average rainfall amount on a 365 days centered moving window (365MA) [mm], (2) the sum of the current and the previous day amounts (2MS) [mm], (3) and (4) two out-of-phase of triangular functions (tr1 and tr2) with frequency 365.25 days, similar to trigonometric coordinates expressing the position of the day in the annual cycle, (5) the dry/wet sequence, i.e. a categorical variable indicating

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the position of a day inside the rainfall pattern (1 = wet, 0 = dry, 2 = solitary wet, 3 = wet day at the beginning or at the end of a wet spell). The first two auxiliary variables are covariates used to force the algorithm to respect more strictly the inter-annual structure and the day-to-day correlation, which are known to exist a priori. The other ones are used to reproduce the dry/wet pattern and the annual seasonality accurately. It has to be remarked that, apart from (3) and (4), which are known deterministic functions imposed as CD, the rest of the auxiliary variables are transformations of the rainfall datum, automatically computed on the TI and co-simulated with the daily rainfall.

To summarize, the main parameters of the algorithm are the following: the maximum scanned TI fraction $F \in (0, 1]$, the search neighborhood radius R , the maximum number of neighbors N , both expressed in number of elements of the time vector, and the distance threshold $T \in (0, 1]$. Recall that, apart from F , each parameter is set independently for each simulated variable. The setup shown in Table 2 is used together with $F = 0.5$ and proposed as a standard for daily rainfall time-series. A sensitivity analysis, not showed here, confirmed the generality of this setup which is not the result of a numerical optimization on a specific dataset, but it is rather in accordance to the criteria used to define the order and extension of the variable time-dependence, as shown below. Applying it to any type of single-station daily rainfall dataset, the user should obtain a reliable simulation without needing to change any parameter or give Supplement. An additional refinement of the setup is also possible, keeping in mind the following general rules:

- R limits the maximum time-lag dependence in the simulation and should be set according to the length of the largest sufficiently repeated structure or frequency in the signal that has to be reproduced. Being interested to condition the simulation upon the inter-annual fluctuations (visible in the 10 years Moving Sum time-series in Fig. 8), we set $R_{365MA} = R_{rainfall} = 5000$ for the 365MA and daily rainfall variables. We recommend keeping R below the half of the training dataset total length, to condition on sufficiently repeated structures only. Regarding dry/wet pattern conditioning, we prefer limiting the variable time dependence within a 21 days

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window ($R_{dw} = 10$). In general R_{dw} should be set between the median and the maximum of the wet spell length distribution, in order to properly catch the continuity of the rainfall events over multiple days.

- N controls the complexity of the conditioning structure but also influences the specific time-lag dependence. For instance, if one increases N, higher-order dependencies are represented, but the weight accorded to a specific neighbor in evaluating the distance between patterns becomes lower. This leads to a less accurate specific time-lag conditioning, but a more complex time-dependence is respected on average. For the rainfall amount and 365MA variables, $N \ll R$ follows the same setup rule as for R_{dw} . In this way, in the initial iterations, the conditioning neighbors will be sparse in a 10 001 days window ($R = 5000$) to respect low-frequency fluctuations, whereas, in the final iterations, they will be contained in a N days window to respect the within-spell variability. The standard value proposed here ($N_{365MA} = N_{rainfall} = 21$) corresponds approximately to the spell distribution median of the Darwin time-series, remaining in the appropriate range for the other considered climates. Conversely, N_{dw} is kept lower in order to focus the conditioning on the small-scale dry/wet pattern. $N_{dw} = 5$ gives in general the best result in terms of dry/wet pattern reproduction, with a gradual degradation of the statistics departing from this value.
- The combination $N = R = 1$ for the 2MS and tr auxiliary variables is equivalent to a lag-(0, 1) dependence and should not be changed since we have no interest in expanding or varying the lag-dependence in this case.
- T determines the tolerance in accepting a pattern. The sensitivity analysis done until now on different types of heterogeneities (Meerschman et al., 2013) confirmed that the optimum generally lies in the interval [0.01, 0.07]. Higher T values usually lead to poorly simulated patterns and lower ones may induce a bias in the global statistics and increase the phenomenon of verbatim copy, i.e. the exact reproduction of an entire portion of data by oversampling the same pattern inside

the TI. For these reasons, we recommend keeping the proposed standard value $T = 0.05$ for all the variables.

- F should be set sufficiently high to have a consistent choice of patterns but a value close to 1, i.e. all the TI is scanned each time, may lower the variability of the simulations and increase the verbatim copy. Using a training dataset representative enough, the optimal value corresponds to a TI fraction containing some repetitions of the lowest-frequency fluctuation that should be reproduced. Considering the randomness of the TI scan, the value $F = 0.5$ chosen in this paper is sufficient to serve the purpose.

3.1 Imposing a trend

The simulation can be constrained to reproduce the same type of trend found in the TI by making use of an auxiliary variable (Chugunova and Hu, 2008). The auxiliary variable proposed here for any type of non-stationarity is $L(y_t) = y_t$, corresponding to the TI time vector. An exact copy $L(x_t) = L(y_t)$ is present in the SG as conditioning data. The parameters for $L(\cdot)$ are set as follows: $R_L = 1$, $N_L = 1$ and $T_L = 0.01$. Therefore, the sampling for each simulated datum $Z(x_t)$ is forced to remain inside the time neighborhood $I(x_t) = y_t \pm T_L V(L)$, $V(L)$ being the total variation of L , i.e. the total length of the series. For example, for the Darwin case, being $V(L) = 26356$ and $T_L = 0.01$, the sampling to simulate $Z(x_t)$ is constrained to $I(x_t) \approx y_t \pm 263$ [days]. In this way, the main statistics are respected, but the local variability is almost completely restricted to the one found inside the training dataset, reproducing the non-stationarity. The following remarks are noteworthy: (i) any type of non-stationarity is automatically imposed by L but, to properly catch the trend and avoid an unnecessary restriction to the local variability, I should be equal to the maximum time interval for which the target variable can be considered stationary; (ii) the simulation cannot be longer than the training dataset, having no basis to extrapolate the trend in the past or future; (iii) the local variability is

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not completely limited by L : a pattern outside the tolerance range (i.e. with a distance over the threshold) could be sampled if no better candidate is found.

3.2 Validation

To validate the proposed technique, the visual comparison of the generated time-series with the reference as well as several groups of statistical indicators are considered. The empirical cumulative probability distributions, computed using the Kaplan–Meier estimate (Kaplan and Meier, 1958), of the daily, the annual and decennial rainfall time-series, obtained by summing up the daily rainfall, are compared using quantile-quantile (qq-) plots. Moreover, the daily rainfall statistics have been analyzed separately for each month considering the average value of the following indicators: the probability of occurrence of a wet day and the mean, standard deviation, minimum and maximum on wet days only. For instance, the standard deviation is computed on the wet days of each month of January, then the average value is taken as representative of that time-series. We therefore obtain a unique value for the reference and a distribution of values for the simulations represented with a box-plot.

Another used validation criterion is the comparison of the dry and wet spells length distributions. Each series is transformed in a binary sequence with zeros corresponding to dry days and ones to the wet days. Then, counting the number of days inside each dry and wet region, we obtain the distributions of dry and wet spells length, that can be compared using qq-plots. This is an important indicator since it determines, for example, the efficiency of the algorithm in reproducing long droughts or wet periods.

Since the DS works by pasting values from the TI to the SG, it is straightforward to keep track of the original location of each value in the training image. If successive values in the TI are also next to each other in the SG, then a patch is identified. A multiple box-plot is then used to represent the number of patches found in each realization as a function of the patch length to evaluate the verbatim copy effect.

The last group of indicators considered is the sample Partial Autocorrelation Function (PACF) (Box and Jenkins, 1976) of the daily, monthly and annual rainfall. Given

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a time-series X_t , the sample PACF is the estimation of the linear correlation index between the datum at time t and the ones at previous time-steps $t-h$, without considering the linear dependence with the in-between observations. For a stationary time-series the sample PACF is expressed as a function of the time-lag h with the following formula:

$$\hat{\rho}(h) = \text{Corr}[X_t - \hat{E}(X_t|\{X_{t-1}, \dots, X_{t-h+1}\}), X_{t-h} - \hat{E}(X_{t-h}|\{X_{t-h+1}, \dots, X_{t-1}\})] \quad (3)$$

where $\hat{E}(X_t|\{X_{t-1}, \dots, X_{t-h+1}\})$ is the best linear predictor knowing the observations $\{X_{t-1}, \dots, X_{t-h+1}\}$. $\rho(h)$ varies in $[0, 1]$, with high values for a highly autocorrelated process. This indicator is widely used in time-series analysis since it gives information about the persistence of the signal. In the case of daily rainfall, the partial autocorrelation is usually very low, while the higher-scale rainfall may present a more important specific time-lag linear dependence. As suggested by Brockwell and Davis (1991), in the absence of any prior knowledge about X_t , an accurate way to detect a significant autocorrelation at a certain lag, is to compare it with an IID $\sim N(0, \sigma^2)$ noise. Such a signal is totally non-autocorrelated and presents a sample PACF ($\hat{\rho}_{AN}$) near zero for any $h > 1$. Moreover, $\hat{\rho}_{AN}$ follows the asymptotic normal distribution $AN(0, n^{-1})$, n being the number of observations in the considered sample. The 95% confidence interval of this distribution can be used to test the significance of any $\hat{\rho}(h)$. That is, in the estimation of the autocorrelation for X_t , at all the $\hat{\rho}(h)$ values within $0 \pm 1.96n^{-1/2}$ can be considered negligible, being of the same magnitude as $\hat{\rho}_{AN}$. Conversely, the values outside these boundaries are probably the expression of a significant autocorrelation and should be reproduced by the simulation. The fact that the variance of $\hat{\rho}_{AN}$ as well as the size of the 95% confidence interval increase with n^{-1} , allows a correct PACF evaluation with respect to the limited information given by the considered sample.

Since the time-series used in this paper are not necessarily stationary, any sample PACF is computed from the standardized signal X_t^S , obtained by applying moving

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average estimation \hat{m}_t and standard deviation \hat{s}_t filters with the following formula:

$$X_t^s = \frac{X_t - \hat{m}_t}{\hat{s}_t}, \quad \hat{m}_t = (2q + 1)^{-1} \sum_{j=-q}^q X_{t+j}, \quad (4)$$

$$\hat{s}_t = [(2q + 1)^{-1} \sum_{j=-q}^q (X_{t+j} - \hat{m}_t)^2]^{-\frac{1}{2}}, \quad q + 1 \leq t \leq n - q$$

where $q = 2555$ (15 years centered moving window). It is important to note that this operation may exclude from the PACF computation a consistent part of the signal ($q + 1 \leq t \leq n - q$), especially on the higher time-scale signal. In the case of the datasets used, the annual time-series is reduced to less than 60 values for Alice Springs and Darwin: a barely sufficient quantity, considering that a generic minimum amount of data for a useful sample PACF estimation given by Box and Jenkins (1976) is of about 50 observations.

4 Results and discussion

To evaluate the proposed technique, a group of 100 realizations of the same length as the reference is generated for each of the 3 considered datasets to obtain a sufficiently stable response in both the average and the extreme behavior. The setup used is the one presented in Sect. 3 with the fixed parameters values shown in Table 2. The obtained results are shown and discussed in the following.

4.1 Visual comparison

Figure 2 shows the comparison between random samples from both the simulated and the reference time-series. For each dataset, the generated rainfall looks similar to the reference: the extreme events inside the 10 years samples are reproduced with an analogous frequency and magnitude. The annual seasonality, particularly pronounced

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in the Darwin series, is accurately simulated as well as the persistence of the rainfall events, visible in the 100 days samples. These aspects are evaluated quantitatively in the following sections.

4.2 Multiple-scale probability distribution

5 The qq-plots of the rainfall empirical distributions are presented in Fig. 3, where all the range of quantiles is considered. The distribution of the daily rainfall (computed on wet days only) is generally respected, although some extremes that are present only once in the reference and, in particular, at the border of the time-series, may not appear in the simulation. It is the case of the Darwin series, with a mismatch of the very upper
10 quantiles. Moreover, the DS being an algorithm based on resampling, the distribution of the simulated values is limited by the range of the training dataset: this is shown in the Alice Springs and Sydney qq-plots, where the distribution of the last quantiles is clearly truncated at the maximum value found in the reference. This result is normally expected using this type of techniques: the direct sampling is of course not able to
15 extrapolate extreme intensities higher than the ones found in the TI at the scale of the simulated signal.

On the contrary, the distribution of the rainfall amount on the solitary wet days is accurately respected, with some realizations including higher extremes than the reference. More importantly, the annual and 10 years rainfall distributions are correctly
20 reproduced and do not show over-dispersion. This phenomenon, common among the classical techniques based on daily-scale conditioning, consists in the scarce representation of the extremes and underestimation of the variance at the higher scale. This problem is avoided here because a variable dependence is considered, up to a 5000 days radius on the 365MA auxiliary variable, that helps respecting the low-
25 frequency fluctuations. We also see that, at this scale, the DS is capable of generating extremes higher than the ones found in the reference, meaning that new patterns have been generated using the same values at the daily scale. This results is purely based on the reproduction of higher-scale patterns: the acceptance threshold value chosen

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a clear upward global tendency which is not present in the superposed randomly-picked DS realization. Generally, the TI is supposed to be stationary or the non-stationarity should be at least described by an auxiliary variable. If it is not the case, as for the Darwin time-series, the algorithm respects the global variation of the reference, but it does not reproduce a specific trend. This problem is treated separately in Sect. 4.6.

4.5 Linear time-dependence

The specific linear time-dependence of the generated and reference signal has been evaluated at different scales using the sample Partial Autocorrelation Function (PACF, Fig. 7, Eq. 3).

At the daily scale, the data show the same level of autocorrelation at lag-1 and a low but significant linear dependence until lag 3 for Alice Springs and Sydney, while Darwin presents a longer tailing which asymptotically approaches the confidence bounds of the Gaussian noise. The DS simulation shows a tendency to a slight underestimation of the lag-1 PACF, with a maximum error around 0.1 (the Sydney time-series). The rest of the lags are reproduced quite accurately.

At the monthly scale, the linear time-dependence structure is clearly related to the annual seasonality, with a negative autocorrelation around lag 6 and a positive one around lag 12. The climate characterization is also evident: from Alice Springs to Darwin we see a more marked seasonality reflected in the PACF. The simulation follows the reference fairly well, with a maximum error around ± 0.1 .

At the annual scale, the limited length of the time-series reduces leads to wider confidence bounds for the non-significant values (see Sect. 3.2). The reference does not show a clear linear time-dependence structure which is not similarly reproduced by the simulation. Some more relevant discrepancy is present in the Darwin series, presenting a more discontinuous structure. However, using such a limited dataset for the time scale considered here, it is difficult to determine if the reference PACF is really indicative of an effective linear dependence.

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4.6 Non-stationary simulation

Figure 8 shows the Darwin time-series simulation realized by imposing the same non-stationarity with the technique proposed in Sect. 3.1. The 10 years moving sum plot shows that the local and global trend of the reference are accurately reproduced: the median of the realizations follows the reference and a variability of about 4 m between the 5th and 95th percentile is present. The accuracy in the global statistics appears to be essentially the same as for the stationary simulation: the only remarkable difference is a modest positive bias in the maximum wet periods length.

The fact that, to impose the trend, the sampling is restricted to a local region of the reference reduces the local variability with respect to the stationary simulation. Consequently, a little increase of the verbatim copy effect occurs.

5 Conclusions

The aim of the paper is to present an alternative daily rainfall simulation technique based on the Direct Sampling algorithm, belonging to multiple-point statistics family. The main principle of the technique is to resample a given dataset using a pattern-similarity rule. Using a random simulation path and a non-fixed pattern dimension, the technique allows imposing a variable time-dependence and reproducing the reference statistics at multiple scales. The proposed setup, suitable for any type of rainfall, includes the simulation of the daily rainfall time-series together with a series of auxiliary variables including: a categorical variable describing the dry/wet pattern, the 2 days moving sum which helps respecting the lag-1 autocorrelation, the 365 days moving average to condition upon inter-annual fluctuations and two coupled theoretical periodic functions describing the annual seasonality. Since all the variables are automatically computed from the rainfall data, no additional information is needed.

The technique has been tested on three different climates of Australia: Alice Springs (desert), Sydney (temperate) and Darwin (tropical savannah). Without changing the

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simulation parameters, the algorithm correctly simulates both the rainfall occurrence structure and amount distribution up to the decennial scale for all the three climates, avoiding the problem of over-dispersion, which often affects daily-rainfall simulation techniques. Being based on resampling, the algorithm can only generate data which are present in the training dataset, but they can be aggregated differently, simulating new extremes in the higher-scale rainfall and dry/wet pattern distributions. The technique is not meant to be used as a tool to explore the uncertainty related to long recurrence-time events, but rather to generate extremely realistic replicates of the datum, to be used as inputs in hydrologic models.

Reproducing a trend in the simulation is also possible by making use of an additional auxiliary variable which simply restricts the sampling to a local portion of the TI. This way, any type of non-stationarity present in the TI is automatically imposed on the simulation. The Darwin example demonstrates the efficiency of this approach in reproducing 100 different realizations showing the same type of trend and global statistics.

In conclusion, the Direct Sampling technique used with the proposed generic setup can produce realistic daily rainfall time-series replicates from different climates without the need of calibration or additional information. The generality and the total automation of the technique makes it a powerful tool for a routine use in scientific and engineering applications.

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Table 1. Summary of the used dataset.

Location	Station	Period [years]	Record length [days]	Missing data [days]
Alice Springs	A.S.Airport	1940–2013	26 347	305
Sydney	S.Observatory Hill	1858–2013	56 662	184
Darwin	D.Airport	1941–2013	26 356	0

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Table 2. Standard setup proposed for rainfall simulation. The parameters are: search window radius R , maximum number of neighbors N and distance threshold T . The variables are: (1) the 365 days Moving Average (365MA), (2) the Moving Sum of the current day and the one before (2MS), (3) and (4) annual seasonality triangular functions (tr1 and tr2), (5) the dry/wet sequence dw and (*) the rainfall amount as the target variable.

Variable	R	N	T	
(1) 365MA	5000	21	0.05	
(2) 2MS	1	1	0.05	
(3) tr1	1	1	0.05	
(4) tr2	1	1	0.05	
(5) dw	10	5	0.05	
(*) rainfall	5000	21	0.05	

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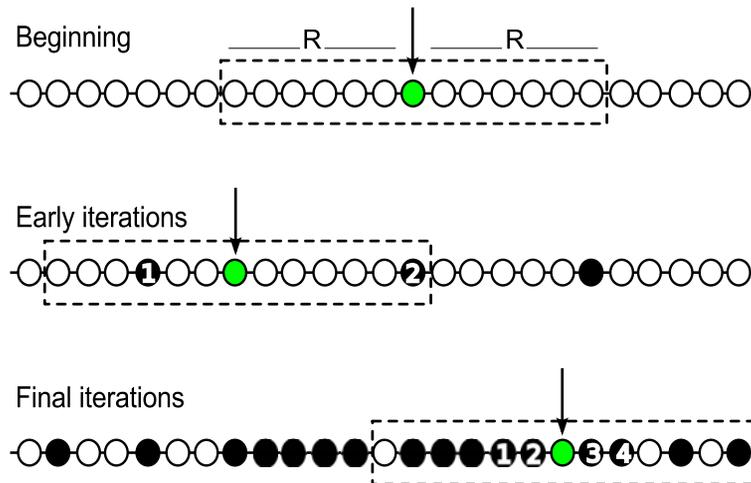


Fig. 1. Sketch of a sequential simulation performed by the Direct Sampling: the chain represents the SG, with circles corresponding to uninformed time-steps and full dots denoting simulated data. The dashed rectangle represents the search neighborhood of radius R , the datum being simulated is indicated with the arrow and the ones composing the data event are numbered. In this example, $R = 6$ and the maximum number of neighbors $N = 4$.

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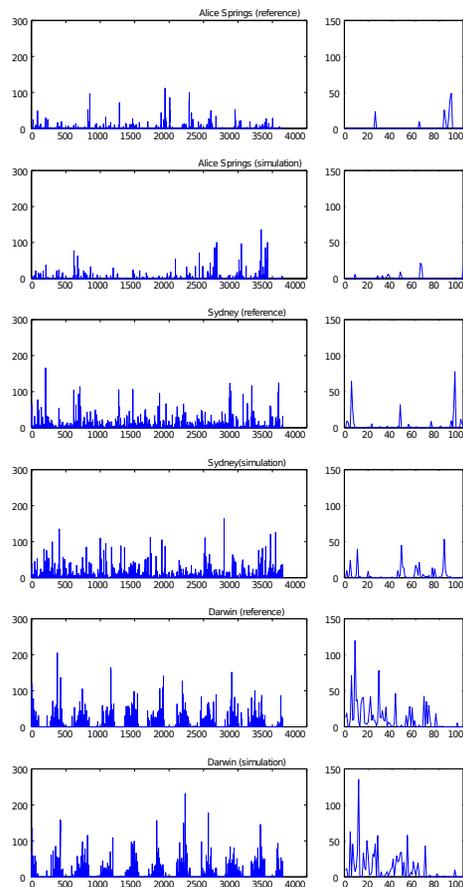


Fig. 2. Visual comparison between the simulated and the reference daily rainfall [mm] time-series: 10 years (left column) and 100 days (right column) random samples.

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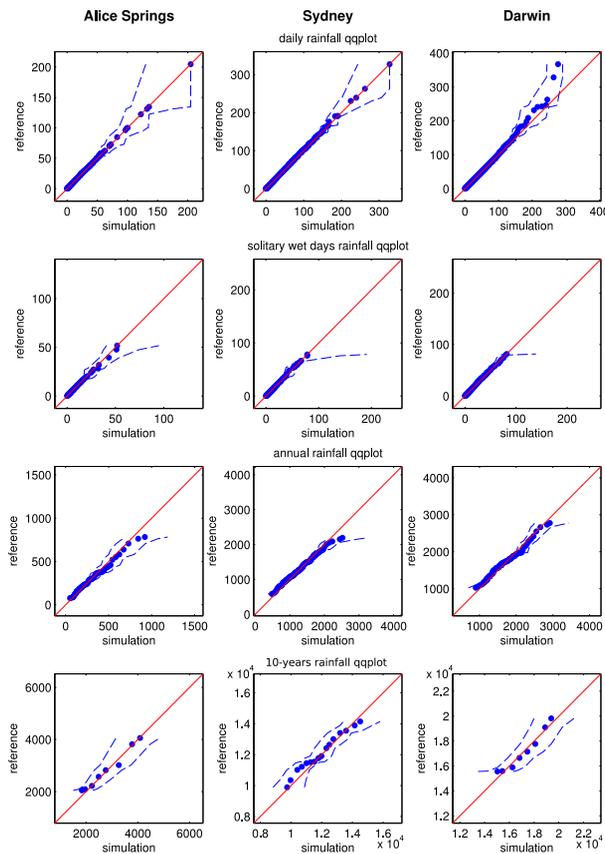


Fig. 3. qq-plots of the empirical probability rainfall amount [mm] distributions: median of the realizations (dotted line), 5th and 95th percentile (dashed lines). The bisector (solid line) indicates the exact quantile match.

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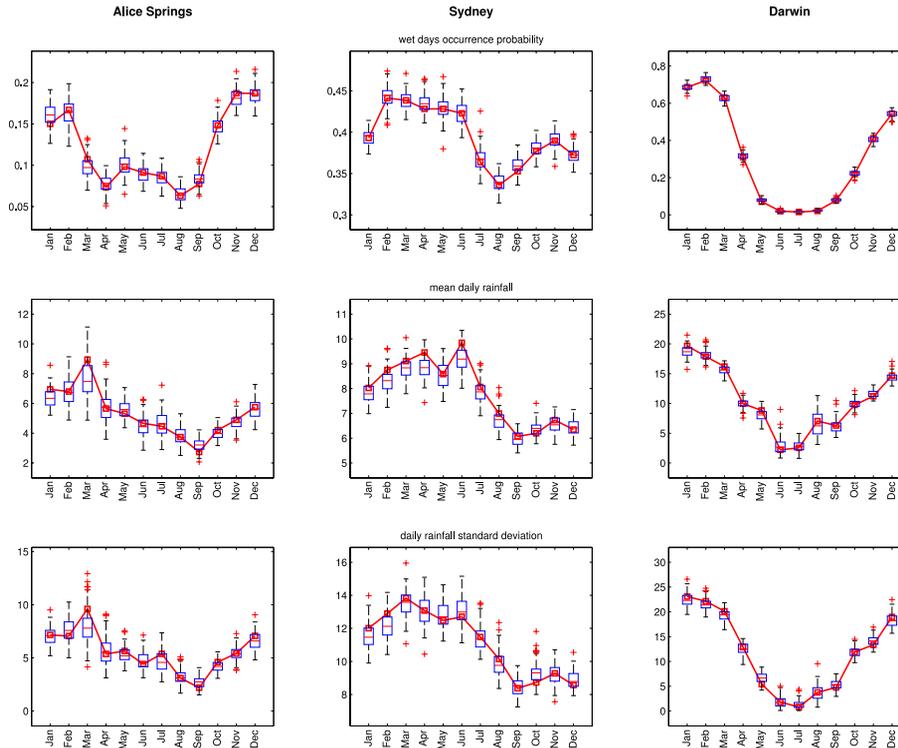


Fig. 4. Box-plots of the average wet days probability, mean daily rainfall amount [mm] and its standard deviation per month. The solid line indicates the reference.

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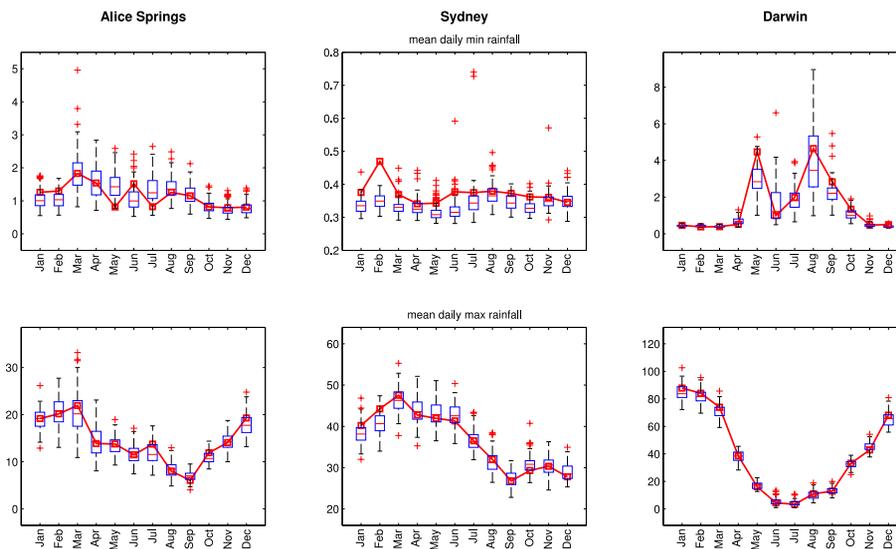


Fig. 5. Box-plots of the average extremes per month. The solid line indicates the reference.

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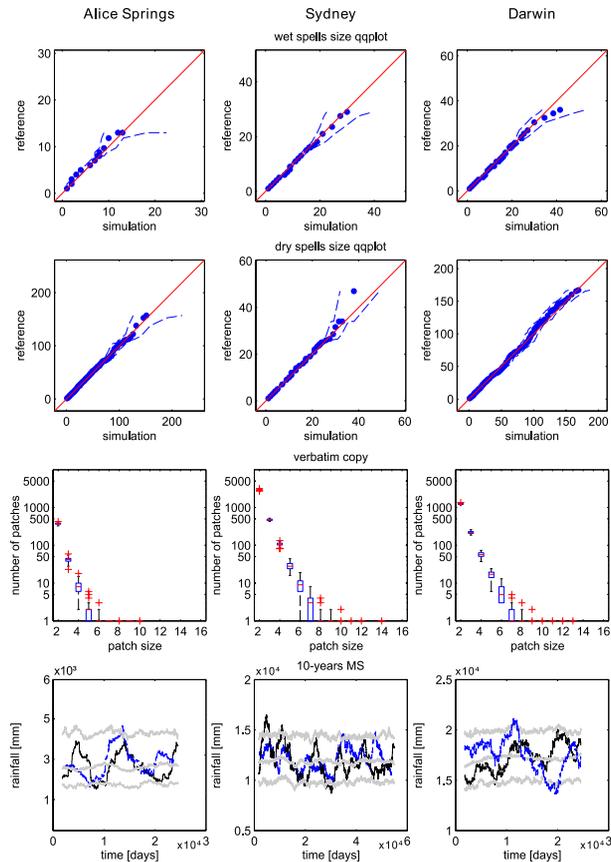


Fig. 6. Main indicators describing the rainfall pattern: qq-plots of the dry and wet spells [days] distributions, verbatim copy box-plots as function of the patch size [days] and daily 10 years Moving Sum (MS) time-series [mm] of the reference (black line), median, 5th and 95th percentile of the realizations (gray lines) and a randomly picked simulation (dashed blue line).

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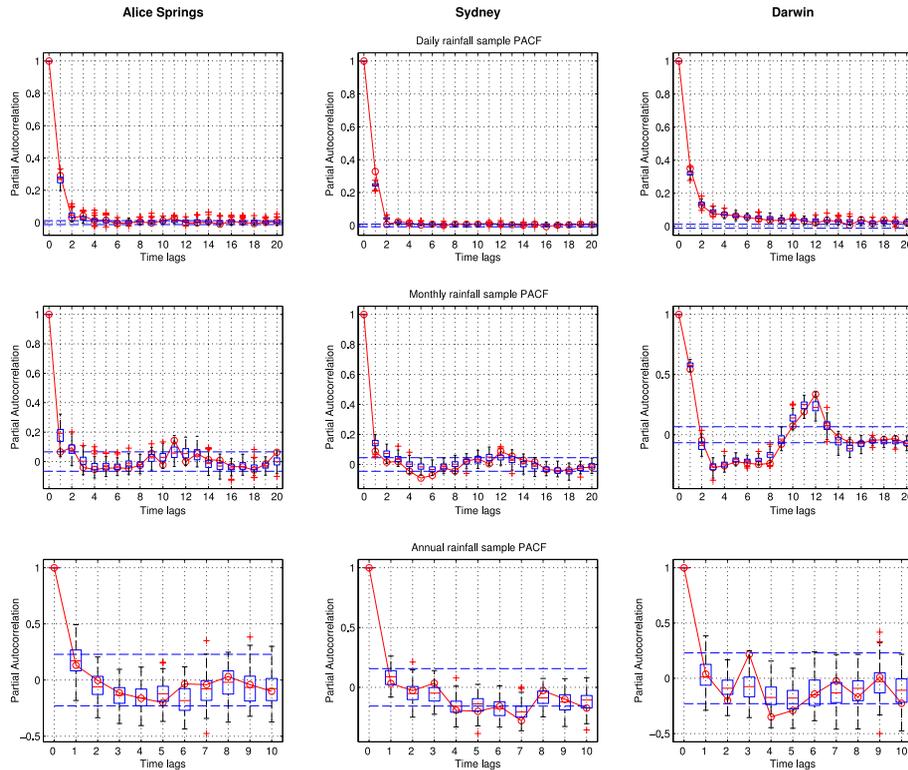


Fig. 7. Sample Partial Autocorrelation Function (PACF) of the daily, monthly and annual rainfall signal: the reference (solid line), 100 DS simulations (box-plots), and confidence bounds for the negligible autocorrelation indexes (dashed lines).

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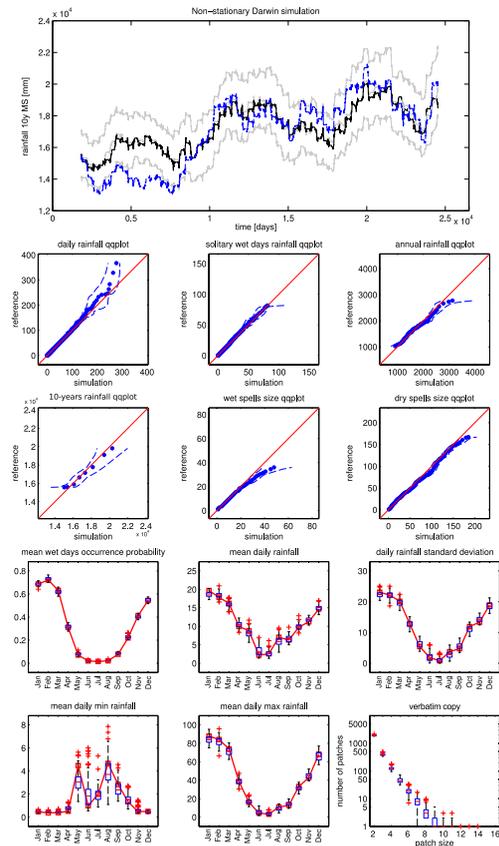


Fig. 8. Darwin daily rainfall non-stationary simulation: 10 years Moving Sum time-series (top) of the reference (black line), median, 5th and 95th percentile of the realizations (gray lines) and a randomly picked simulation (dashed blue line); main quantile-comparisons (center); main seasonal indicators and verbatim copy box-plot (bottom).

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