We would like to thank Dr Serinaldi and the anonymous reviewer for their review of the paper “Non-stationarity in annual maxima rainfall across Australia-implications for Intensity-Frequency-Duration (IFD) relationships”. We have considered the Reviewers’ comments and provided detailed descriptions of how each comment will be addressed in the revised manuscript below:

Response to Reviewer 1 (Dr Serinaldi)

Specific comments

Dr Serinaldi’s specific comments center around three main themes. The first is the use of change point analysis to test for non-stationarity in a data series, the second is the application of the CUSUM test (and identifying multiple change points) and the third is the use of the term “regime shift” in preference to “non-stationarity” when discussing variability in the annual maxima rainfall timeseries. Each of these issues were further built on in the technical remarks provided by Dr Serinaldi and therefore are addressed in detail below.

Technical remarks

1. Please, consider to check the significance of (bias-corrected) serial correlation (if this was not done) because it can affect the results of change point analyses (see e.g. Serinaldi and Kilsby (2015a) and references therein for a discussion on Mann-Kendall and Pettitt, which however holds true also for e.g. CUSUM and similar). As shown above, apparent regime shifts can be artifacts resulting from hidden persistence.

Response: In response to the Dr Serinaldi’s suggestion, the Durbin-Watson (DW) statistic was used to test for autocorrelation (serial correlation) in the annual maxima timeseries (Durbin and Watson (1950, 1951)). The Durbin-Watson statistic tests the null hypothesis that the residuals from an ordinary least-squares regression are not autocorrelated against the alternative that the residuals follow an AR1 process. The Durbin-Watson statistic ranges in value from 0 to 4. A value near 2 indicates non-autocorrelation; a value toward 0 indicates
positive autocorrelation; a value toward 4 indicates negative autocorrelation. Typically, tabulated bounds are used to test the hypothesis of zero autocorrelation against the alternative of positive first-order autocorrelation. For the sample size in our case (~100) and a linear trend model with intercept the $d_{\text{lower}} = 1.522$ and $d_{\text{upper}} = 1.562$ for 1% significance.

All DW statistic values were found to be greater than the 1.562 (the upper bound for 1% significance) providing no evidence to reject the null hypothesis (see figure 1 showing the distribution of all DW statistic values for the 1-day annual maxima timeseries at each site).

A discussion of the DW test results are included in the revised paper (Section 2.2) to demonstrate that the annual maxima data does not suffer from serial correlation and therefore the statistical tests used in the change point analysis is appropriate.

Figure 1: Box plot of Durbin-Watson statistic for the 1-day annual maxima timeseries at 96 stations (red line indicates $d_{\text{upper}} = 1.562$ for 1% significance)

2. P3453L20-25: In my opinion, such lines reflect some confusion on this topic. Trends or change points in finite time series do not imply nonstationarity. Nonstationarity cannot be in principle significant or not significant, because it is an assumption made on the underlying process that can be introduced only if we know the underlying nonstationary dynamics (physical equations, well-defined changes with a clear cause such as flow regime changes due to dams operation, etc.). Please consider to reword
Response: Dr Serinaldi’s review has highlighted some important points regarding the use of the term “stationarity” and if/where it is applicable in our study. In particular he questioned whether the term “regime shifts” was more fitting in describing our findings of change points in the annual maxima rainfall timeseries that are possibly attributable to climate shifts. Dr Serinaldi stated that (following Koutsoyiannis and Montanari, 2014; Serinaldi and Kilsby, 2015) stationarity is a concept referring to models rather than to timeseries. In our case the model is the IFD curve. Thus the text that describes the assumptions in the IFD development of stationarity in the underlying processes (i.e. the statistical properties of the rainfall do not change over time and that the chance of an extreme event occurring is the same at any point in time (past or future)) is relevant, however we agree that the text discussing change points in the rainfall data may be misleading where the term non-stationarity has also been used. However it is interesting to note that many studies have also used the word stationarity and non-stationarity when describing similar timeseries (e.g. Ishak et al. 2013, Westra and Sisson 2011, Wagesho et al 2013, Wilby 1998, etc), therefore there appears to be widespread disagreement on the use of this term. Despite this, on further review of the journal papers provided by the Reviewer we agree that in our case the term “regime shift” is more suitable. Given the above we have revised the text in the paper to reflect this. In particular, we have use the term non-stationarity only when referring to the IFD development, however the sections of the paper that are focused on identifying change points in the rainfall timeseries have been edited and the term “regime shift” has been used in preference. Further, as per the reviewer’s suggestion (in his specific comments) we also changed the title of the paper to “Regime shifts in annual maxima rainfall across Australia – implications for Intensity–
Frequency–Duration (IFD) relationships”. The associated text also includes the references provided by the reviewer.

3. P3457L25: Please consider to reword, e.g. “LP3 was not rejected at x% significance level for all series (or n series out of N)”.

Response: The sentence in question was reworded as suggested to read “Here the null hypothesis is that the data fits the Log-Pearson III distribution (the alternate is that the data does not follow the Log Pearson III distribution). All p-values were greater than 0.05 (average p-value was 0.75), for all series (30min to 72hr durations at Brisbane, Sydney and Melbourne), therefore we accept the null hypothesis at the 5% significance level.

4. P3458L12-15: I do not know AR&R, but it is not clear to me why return periods defined on annual maxima should be adjusted for PDS. Usually we do the opposite when we start from PDS and we need the actual AMAX return periods (under suitable conditions such as Poisson arrival dynamics, etc.). Please clarify.

Response: The methodology adopted in this paper to calculate return periods of annual maxima specifically follows that outlined in Australian Rainfall and Runoff (1987), Engineers Australia’s guide to estimating and utilising IFD information. Published IFD currently used by industry in Australia are based on this method. The updated IFD (which are NOT currently used in operation) are based on a revised statistical methods (for example, a Generalised Extreme Value (GEV) frequency distribution was fitted to the annual maxima rather than Log Pearson III and extension of sub-daily rainfall statistics to daily read stations is conducted with Bayesian Generalised Least Squares Regression rather than PCA). The purpose of adopting the AR&R 1987 method was to assess the implications of varying data lengths and climatic variability on the resulting IFD (which have been historically used and are currently still in use) and to highlight the issue of underlying variability in the annual...
maxima that should be appropriately considered and addressed in the current (and future) revision of the IFD estimates.

5. **P3459L9: As mentioned above, step changes and nonstationarity are very different concepts and surely not synonyms.**

Response: This has been revised and clarified as per discussion above (Comment 2).

6. **P3459L17-24: Leaving aside the use of the term nonstationarity, CUSUM identifies automatically the change point location and does not split the time series in two halves. If the Authors mean that the test proceeds based on subsequent dyadic partitions, this is right, but for such short time series it is actually quite difficult (and not meaningful) to go beyond 2-4 changes. Please note that many other refined techniques are available for segmentation... of course, a question rises about the (physical) meaning of such refined segmentations...**

Response: This has been clarified in the text. The data is not split in equal halves for the CUSUM test, it is split into two portions, which may or may not be equal. However, unless a moving window is used (say 20 years as we did for the Mann-Whitney, multiple regime shifts could still be missed using this method.

7. **P3459L25-P3460L10: Following the previous remark, my interpretation of P3473Fig5 is a bit different. The almost uniform spread of changes across the decades denotes that such changes occur quite randomly, and sincerely I cannot see a tendency to cluster in the east coast. We may see something in panel (b), but the spatial distribution of the stations is not uniform and we cannot exclude that such stations are spatially correlated, as they are subject to similar climate forcings (thus reducing the evidence for changes). Note that spatial correlation is another factor that can strongly affect the outcome of such a type of tests (see e.g. Douglas et al. (2000), Yue et al. (2003), Guerreiro et al. (2014), among others)**

Response: We agree with the Reviewer that there is almost a uniform spread of changes across the decades based on panel (a). We state in the paper that “the large-scale climate phenomena impact various regions of Australia at different times of the year and to varying
degrees, therefore it is not surprising that the timing of shifts in the annual maxima timeseries varies spatially and temporally.” However has been further clarified in the revised paper.

The clustering along the east coast can only be clearly seen in panel B. The text has been revised to clarify this.

The spatial correlation of the annual maxima timeseries was investigated as per the reviewer’s suggestion. We found that less than 9% of all possible pairings of rainfall data sets display a significant (yet weak) correlation at the 5% level ($r > 0.2$, significance based on $n=100$). Only 8 pairings (out of 4465) were correlated at 0.5 or higher. It was also found that stations located more than 500km apart were unlikely to be correlated and that the strength of the correlation reduced as distance increased between the pairs (see Figure 2).
8. Section 3.2: Again, my interpretation of P3474Fig6 and P3475Fig7 is a bit different. If I’m right, box plots for IPO(-) summarize the distribution of 41 AMAX (1913-1920 and 1945-1977), while we have 67 AMAX for IPO(+) box plots. For such sample sizes, inferring difference in distribution based on box plots is a bit hard (at least). My suggestion is to use some formal two-sample goodness-of-fit tests such as the two-sample Kolmogorov-Smirnov or similar, thus accounting for sampling uncertainty and different sample sizes. In any case, comparing box plots (overlooking the large uncertainty of the quantile estimates) is not informative and does not provide a quantitative assessment, especially in this case where differences between IPO(-) and IPO(+) regimes are really hard to recognize.

Response: As suggested a two sample KS test was applied to the data (to test the significance of the difference between the two IPO distributions). It was found only the results for the Sydney were statistically significant (p-value <0.1). A discussion of the two-sample KS test have been included in the revised paper as suggested. Given the results of the significance test we only further investigate the impact of the IPO step induced regime shifts on the IFD for Sydney (rather than all three stations).

9. The same holds for P3475Fig7: if I’m right, this diagram shows the differences Δ (in %) between the point estimates of rainfall return levels obtained by LP5 distributions fitted on 41 and 67 AMAX. It is almost superfluous to highlight how large the uncertainty of such a point estimates can be. I suggest a fairer check based on a simple bootstrap procedure. For each duration:
1. resample with replacement IPO(-) and IPO(+) time series to obtain two new B-samples;
2. for each B-sample refit LP3, compute the required LP3 return levels and calculate the difference $\Delta^{(B)}$ as for the observed data;
3. repeat previous steps B times (e.g. 1000) and store the obtained B differences (for each ARI). These values can be used to build the empirical distribution of the differences $\Delta^{(B)}_i$, $i = 1, \ldots, B$. This distribution describes the effects of sampling and parameter estimation uncertainties under the hypothesis of existence of two different regimes;
4. Use the B $\Delta^{(B)}$ values to build confidence intervals (CIs) at a given confidence level (e.g. 95%). If these CIs include $\Delta = 0$, then there is not evidence for a significant difference, otherwise we can conclude the opposite.

I think this is a better way to provide a quantitative assessment. Of course, conclusions concern the effects of possible regime shifts and not of nonstationarity. Section 3.2 should be reworded according to the results of the analyses suggested above.

Response: We would like to thank Dr Serinaldi for this suggestion and we have completed the analysis as per the steps outlined above. This has resulted in a more robust test of the effects of the IPO induced regime shifts. As per the previous comment, we only apply this method to the Sydney station (since this was the only station where the KS statistic was significant). Figure 7 has been replaced by the figure below that shows the percentage difference in the rainfall intensity estimate between the IPO positive and negative phases. Positive (negative) values represent an increase (decrease) in rainfall intensity during IPO positive compared to IPO negative. The relationships observed are robust for most durations and return intervals given that the CIs do not include 0 (other than 30mins 5 years, 2 hours 50 years, 6 hours 20 years and 24 hours 5 years).
10. Section 4: as for Section 3.2, this section should be reworded according to the updated results.

Response: Noted, this has been revised.

11. Please avoid sentences such as that in P3464L27-29 and P3465L1-3: even after more accurate analyses, there is no way to make unquestionable conclusions about nonstationarity if we do not identify a well-defined mechanism of evolution which is almost perfectly predictable (at least, at the time scales of interest).

Response: The sentence “Based on the results of this study, and literature cited within this paper, we emphasise that there undoubtedly is non-stationarity in historical short duration rainfall extremes but the characteristics and causes of this non-stationarity vary from location to location and decade to decade – something which must be considered and accounted for when attempting to estimate IFD design rainfalls and prior to quantifying...”
This study has highlighted the existence of regime shifts in annual maxima rainfall data in Australia. The driving mechanisms of these regime shifts are likely to vary from location to location and decade to decade. However, these shifts are typical of many natural phenomena and can be described by processes characterized by long range dependence (or regime-switching processes) and captured by hidden Markov models (or similar), resulting in a mixture of distributions that alternate stochastically according to the transition probability from one regime to the next (e.g. Serinaldi and Kilsby, 2015a).

While the strategy of defining IFDs for two (or more) different regimes (e.g. Serinaldi and Kilsby (2015a)) currently only partially solves the problem, as we often do not know the beginning or the end of a specific regime (be it rainfall or climate driver), recent work has focused on optimizing designs and planning strategies based on the range of what is plausible rather than a reliance on knowing the current and future climate state (e.g. Mortazavi-Naeini et al., 2015). At the same time, work is also underway on seamless prediction at a range of timescales and if/when this eventuates the results discussed here become even more important/useful. Nevertheless, the immediate usefulness of the insights presented here occurs when first establishing the IFD, as an approach similar to that employed here can be used to determine if the underlying data are biased to a mostly wet or mostly dry regime (or a mix of both) which then provides an indication as to whether the IFD is likely to be an over- or underestimate of the true risk. Importantly, this issue needs to be considered and accounted for when attempting to estimate IFD design rainfalls and prior to quantifying how those IFD estimates might change in both the near and long-term future.”
Response to Reviewer 2 (Anonymous reviewer)

Specific comments:

1. Recent studies (Montanari and Koutsoyiannis, WRR, 2014; and references therein, Koutsoyiannis, JH, 2006) show that modeling approaches which consider non-stationarity of real world time series without examining the properties of the stochastic processes, may be inappropriate. The way the authors test and claim for non-stationarity in the extremes is inadequate and quite limited. In fact, it has been shown in some studies (Serinaldi and Kilsby, AWR, 2015) that non-stationary models may increase the uncertainties and that traditional concepts should still be retained as benchmarks. Thus, the authors’ skepticism about the BoM and ARR’s existing approaches may not be justified.

Response: Based on this review and the review of our paper provided by Dr Serinaldi we have improved the paper as follows:

a. We have included a test for serial correlation (Section 2.2) using the Durbin-Watson (DW) statistic. The Durbin-Watson statistic tests the null hypothesis that the residuals from an ordinary least-squares regression are not autocorrelated against the alternative that the residuals follow an AR1 process. All DW statistic values were found to be greater than the 1.562 (the upper bound for 1% significance) providing no evidence to reject the null hypothesis (see figure below showing the distribution of all DW statistic values for the 1-day annual maxima timeseries at each site).

b. We have addressed the issue of potential spatial correlation among rainfall sites. We found that less than 9% of all possible pairings of rainfall data sets display a significant (yet weak) correlation at the 5% level (r >0.2, significance based on n=100). Only 8 pairings (out of 4465) were correlated at 0.5 or higher. It was also found that stations located more than 500km apart were unlikely to be correlated and that the strength of the correlation reduced as distance increased between the pairs. This is not surprising given annual maximum rainfall events are due to synoptic scale processes. This is include in revised Section 2.2 of the paper.

c. We have revised the text in the paper with respect to the use of the term non-stationarity. In particular, we use the term non-stationarity only when referring to the IFD development (which is deemed appropriate given the IFD is essentially a model), however the sections of the paper that are focused on identifying change points in the rainfall timeseries have been edited and the term “regime shift” has been used in
preference. Further, we have also change the title of the paper to “Regime-shifts in
annual maxima rainfall across Australia – implications for Intensity–Frequency–
Duration (IFD) relationships”. The associated text also includes the references
provided by the reviewer

2. Also, the definition of return period itself (and equivalently that of ARI) may change
in the non-stationary setting (Salas and Obeysekara, ASCE JHE, 2013). Moreover,
although the title mentions “implications for Intensity–Frequency–Duration (IFD)
relationships”, this paper only presents a discussion (Section 3.2) which contains
rather generic discussion on how non-stationarity may affect such relationships,
without carrying out any analysis on how the observed-period IFD relationships
actually change because of non-stationarity (such as that done by Cheng and
AghaKouchak, Sc. Re- ports, 2014), bringing into question the novelty, utility and
scientific contribution of this study.

Response: We agree with the reviewer that the ARI may change in the non-stationary
setting. In fact, that is the point we are making in our paper; that depending on when the
data is sampled from to generate the IFD, it may be biased to either a wet or a dry phase
(or surplus or absence of high intensity events) and therefore would have consequences
on the resulting return period for individual rainfall depths. Indeed we suggest in our
discussion “that a separate set of IFDs could be developed for use in high risk modelling
for engineers who need to account for the ‘worst case’ (in a similar manner to climate
change allowances). This second set of IFDs could be developed based on the periods of
elevated annual maxima alone (for those stations with clearly defined epochs of annual
maxima) such that if we were to enter such an epoch, designs based on these estimates
would be robust for the duration of such a period.” We disagree with the reviewer that we
do not show how nonstationarity (which we will now term regime shifts) may affect the
IFD relationships. This is demonstrated in Section 3.2 “Effect of non-stationarity on IFD
estimation” where we recalculate the IFD curves for the difference phases of the IPO.
Importantly we show that the return period is different for the various rainfall depths and
durations depending on the underlying rainfall dataset (i.e. depending on whether it is
3. The authors mention, in their conclusions, "The research presented here demonstrates that information currently available on natural variability...can act as a guide to the baseline..." - this is a fat-fetched conclusion. The present research, however, doesn’t provide any guidelines on how this baseline can be defined.

Response: This has been clarified in the revised paper. Our intention here was to emphasize that, for regions where large-scale climate drivers operate on a multi-year to multi-decadal timescales and are known to influence extreme rainfall events, we can use this information to determine if the climate statistics on which the IFD are based are likely to be biased or missing crucial information.

4. Is IPO the same as PDO (Pacific Decadal Oscillations)? If it was known apriori that locations such as Melbourne are not affected by the IPO, why was it chosen for the analysis? Perhaps a more appropriate approach would consider several natural variability modes, as well as forced changes and investigate their individual effects on rainfall extremes.

Response: The IPO is not the same as the PDO. The IPO is a Pacific Basin wide phenomena rather than just the north Pacific that is represented by the PDO. There are similarities between the two timeseries however and they are significantly correlated. According to Salinger et al 2001 “The IPO may be a Pacific-wide manifestation of the PDO, excluding subdecadal time scales, and seems to be part of a continuous spectrum of low frequency modulation of ENSO, and so may be partly stochastic”.

It is true that some existing studies suggest that the IPO signal on rainfall tends to be weaker in Melbourne due to competing influences from the Southern Ocean, however we cannot say that Melbourne rainfall/climate is “not affected by the IPO”. Some studies suggest IPO significantly effects rainfall characteristic in Melbourne (e.g. Verdon et al 2004, Gallant et al 2012) while others do not, therefore in our study we do not make any assumptions about IPO effects on rainfall maxima in Melbourne and include it in our investigation. Our results suggest there is a relationship where “all events (other than 72
hours) with a 2-year ARI are associated with a higher rainfall intensity estimate in IPO positive for Melbourne, however the reverse is true for the less frequent events.”

While we agree with the reviewer that there are several modes of natural climate variability that may have an effect the extreme rainfall from year to year, in our study we were specifically interested in climate drivers that are likely to force a regime shift in extreme rainfall (similar to that observed for flood risk). Therefore we were interested in drivers that operate on a decadal to multi-decadal timescale (as is the case for the IPO). Other drivers (such as ENSO, Indian Ocean Dipole, Southern Annular Mode) tend to influence rainfall in Australia on much shorter timescales. However, if this method was to be applied to regions other than east coast Australia (i.e. where IPO is known not to be the primary driver on decadal to multi-decadal timescales), other potential sources of decadal to multi-decadal variability would need to be identified.

5. GEV distribution is usually deemed appropriate for annual maxima. How does the GEV distribution fit the data at hand? How are spatial dependence between extremes taken into account? Why not consider peak-over-threshold approach?

Response: The Reviewer’s point is correct, the GEV distribution does indeed fit the annual maxima data well. In fact the updated IFD (which are NOT currently used in operation) are based on a revised statistical methods that includes fitting the GEV distribution to the data in preference to the Log Pearson III. However, the methodology adopted in this paper (including fitting the Log Pearson III) to calculate return periods of annual maxima deliberately follows that outlined in Australian Rainfall and Runoff (1987), Engineers Australia’s guide to estimating and utilising IFD information. The purpose of adopting the AR&R 1987 method was to assess the implications of varying data lengths and climatic variability on the resulting IFD (which have been historically used and are currently still in use) and to highlight the issue of underlying variability in
the annual maxima that should be appropriately considered and addressed in current and
future revisions of the IFD estimates. Further, as an additional check the KS goodness of
fit test was applied to test if the Log Pearson III was a reasonable fit to the data. Here the
null hypothesis is that the data fits the Log-Pearson III distribution (the alternate is that the
data does not follow the Log Pearson III distribution). All p-values were greater than 0.05
(average p-value was 0.75), for all series (30min to 72hr durations at Brisbane, Sydney
and Melbourne), therefore we accept the null hypothesis at the 5% significance level.

This has been clarified in the revised paper.

6. Claims such as “we emphasize that there undoubtedly is non-stationarity in historical short duration rainfall extremes” might be inappropriate for reasons stated above

Response: We agree and this has been revised and the discussion extended. The following
text replaces the sentence above:

“This study has highlighted the existence of regime shifts in annual maxima rainfall data
in Australia. The driving mechanisms of these regime shifts are likely to vary from
location to location and decade to decade. However, these shifts are typical of many
natural phenomena and can be described by processes characterized by long range
dependence (or regime-switching processes) and captured by hidden Markov models (or
similar), resulting in a mixture of distributions that alternate stochastically according to
the transition probability from one regime to the next (e.g. Serinaldi and Kilsby, 2015a).
While the strategy of defining IFDs for two (or more) different regimes (e.g Serinaldi
and Kilsby (2015a)) currently only partially solves the problem, as we often do not know
the beginning or the end of a specific regime (be it rainfall or climate driver), recent work
has focused on optimizing designs and planning strategies based on the range of what is
plausible rather than a reliance on knowing the current and future climate state (e.g.
Mortazavi-Naeini et al., 2015). At the same time, work is also underway on seamless
prediction at a range of timescales and if/when this eventuates the results discussed here become even more important/useful. Nevertheless, the immediate usefulness of the insights presented here occurs when first establishing the IFD, as an approach similar to that employed here can be used to determine if the underlying data are biased to a mostly wet or mostly dry regime (or a mix of both) which then provides an indication as to whether the IFD is likely to be an over- or underestimate of the true risk. Importantly, this issue needs to be considered and accounted for when attempting to estimate IFD design rainfalls and prior to quantifying how those IFD estimates might change in both the near and long-term future.”

7. Literature review pertains mostly to studies on Australian datasets, whereas much work on similar ideas are also carried out elsewhere.

Response: The literature review has been extended to include the references mentioned by the Reviewer as well as the following international papers:

Regime shifts in Annual Maxima Rainfall across Australia– Implications for Intensity-Frequency-Duration (IFD) relationships

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Abstract

Rainfall Intensity-Frequency-Duration (IFD) relationships are commonly required for the design and planning of water supply and management systems around the world. Currently, IFD information is based on the ‘stationary climate assumption’ - that weather at any point in time will vary randomly and that the underlying climate statistics (including both averages and extremes) will remain constant irrespective of the period of record. However, the validity of this assumption has been questioned over the last 15 years, particularly in Australia, following an improved understanding of the significant impact of climate variability and change occurring on interannual to multidecadal timescales. This paper provides evidence of regime shifts in annual maxima rainfall timeseries using 96 daily rainfall stations and 66 sub-daily rainfall stations across Australia. Further, the effect of these regime shifts on the resulting IFD estimates are explored for three long-term sub-daily rainfall records (Brisbane, Sydney and Melbourne) utilising insights into multidecadal climate variability. It is demonstrated that IFD relationships may under- or over-estimate the design rainfall depending on the length and time period spanned by the rainfall data used to develop the IFD information. It is recommended that regime shifts in annual maxima rainfall be explicitly considered and appropriately treated in the ongoing revisions of Engineers Australia’s guide to estimating and utilising IFD information, “Australian Rainfall and Runoff”, and that clear guidance needs to be provided on how to deal with the issue of regime shifts in extreme events (irrespective of whether this is due to natural or anthropogenic climate change). The findings of our study also have important implications for other regions of the world that exhibit considerable hydroclimatic variability and where IFD information is based on relatively short data sets.
1. Introduction

Information on rainfall event intensity, frequency and duration (IFD, or IDF as it is known in some countries) plays a critical role in the design of dams, bridges, stormwater drainage systems and floodplain management. Dependent upon the application, information is required for event-durations ranging from hours to several days. The development of IFD relationships were first proposed by Bernard (1932) and since then different versions of this relationship have been developed and applied worldwide (e.g. Bara et al. 2009, Chen 1983, Hershfield 1961, IHP-VII 2008, Nhat et al. 2006, Raiford et al. 2007).

Historically, in Australia, IFD design rainfall curves were developed by the Australian Bureau of Meteorology (BoM) for durations ranging from 5 minutes to 72 hours and Average Return Intervals (ARI) of 1 year to 100 years (however, recently additional durations and ARIs have also been developed). Up until very recently IFD information available to (and used by) engineers and hydrologists were developed 25 years ago, as part of Engineers Australia publication Australian Rainfall and Runoff (AR&R) in 1987. New IFD information was released early in 2013 after a major revision of IFD information carried out by Engineers Australia. Importantly, the revised IFD information is based on a longer and more extensive rainfall data set (http://www.bom.gov.au/water/designRainfalls/ifd/). However, the BoM and Engineers Australia still recommend to use the AR&R 1987 information for existing flood studies and the probabilistic rational method and to conduct sensitivity testing with the revised 2013 AR&R parameters including the new IFD design rainfalls (http://www.bom.gov.au/water/designRainfalls/ifd/index.shtml).

At the time of writing, the revised IFD information does not take into account the impact of climate change on IFD estimates. This is part of ongoing research commissioned through Engineers Australia. It is also not yet clear how or if the role of natural climate variability is going to be considered. Of concern is the fact that currently, estimates of IFD are based on
the assumption that “climatic trend, if it exists in a region, has negligible effect on the design
intensities” (Pilgrim 1987). This is known as the ‘stationary climate assumption’ (i.e.
the statistical properties of the rainfall do not change over time) and implies that the chance of
an extreme event occurring is the same at any point in time (past or future). However, the
validity of this assumption has been questioned over the last 15 years following
demonstration of the significant impact of climate variability occurring on interannual to
multidecadal timescales in Australia. For example, research has shown that annual maximum
flood risk estimates in Australia vary depending on climate state (e.g. Ishak et al. 2013, Kiem
et al. 2003, Leonard et al. 2008). Importantly these studies demonstrate that founding flood
risk estimates on an unsuitable time period has the potential to significantly underestimate (or
overestimate) the true risks. This may apply to design rainfall also given that current IFD
estimates are based on varying lengths of data spanning different time periods (the latest IFD
estimates are based on all daily-read stations with 30 or more years of record and all
continuously-recording stations with more than 8 years of record).

Khaliq et al. (2006) explained that the traditional idea of probability of exceedance and return
period are no longer valid under non-stationarity. Recently, Jakob et al (2011a) found that
rainfall quantile estimates derived for Sydney Observatory Hill for the period 1976 to 2005
show significant decreases across durations from 6 minutes to 72 hours. Jakob et al (2011b)
subsequently extended the sub-daily rainfall data analysis to 31 sites located in southeast
Australia, assessing variations in frequency and magnitude of intense rainfall events across
durations from 6 minutes to 72 hours. This study identified two different trends in the data
sets, a decreasing trend in frequency of events at durations of 1-hour and longer for sites in
the north of the study region, while sites in the south cluster displayed an increase in
frequency of events, particularly for sub-hourly durations. Importantly Jakob (2011a, 2001b)
concluded that, for at least some regions of Australia, trends found in historical records has
the potential to significantly affect design rainfall estimates. Westra and Sisson (2011) also investigated evidence of trends in extreme precipitation at sub-daily and daily timescales (1965-2005) using a spatial extreme value model. They identified a statistically significant increasing trend in precipitation extremes for the sub-daily data set, however at the daily timescale no change in annual maximum rainfall could be detected with the exception of southwest Western Australia (Westra and Sisson 2011). Further, Yilmaz and Perera (2014) conducted change point analysis for extreme rainfall data for storm durations ranging from 6 minutes to 72 hours in Melbourne, and found evidence of regime shifts, concluding the year 1966 as a statistically significant change point. Yilmaz et al (2014) then investigated changes in extreme rainfall through trend analysis, non-stationarity tests and non-stationary GPD models (NSGPD) for Melbourne. They found statistically significant extreme rainfall trends for storm durations of 30 minutes, 3 hours and 48 hours, however for above storm durations there was no evidence of a regime shift (which they termed 'non-stationarity') according to statistical non-stationarity tests and non-stationary GPD (Yilmaz et al 2014).

A limitation of the analysis presented by Westra and Sisson (2011) and Jakob et al (2011a, 2011b) is that they tested for linear trends in the rainfall data series based on the hypothesis that extreme rainfall events would have either decreased, increased or exhibited no trend over the time period being investigated. However these are not the only attributes of trend detection, since annual rainfall maxima may also cycle through interannual to multidecadal periods (note that Westra and Sisson (2011) also investigated possible links between extreme rainfall and annual fluctuations in the El Niño/Southern Oscillation (ENSO)). Therefore, depending on what time period the annual rainfall maxima data are derived from (in reference to any long term cycles or epochs) the observed trends may be misleading or even not apparent (leading to the misconception that regimes shifts are non-significant or not an important consideration). Recently Yilmaz et al (2014) investigated the potential impact of...
the Interdecadal Pacific Oscillation (IPO) on extreme rainfall and resulting IFD for a case study in Melbourne. They concluded that, the IPO negative phase can be the driver of higher rainfall intensities for long durations and high return periods. However, the trends in extreme rainfall data and differences in rainfall intensities for short storm durations and return periods could not be explained with the IPO influence. Given that Melbourne is located in south-east Australia, where the influence of the IPO is temporally variable due to other climate drivers operating (acting to enhance or suppress impacts, see Kiem and Verdon-Kidd 2010; 2009), the research by Yilmaz et al (2014) provides promise for developing relationships between extreme rainfall and the IPO for regions where the IPO may have a more consistent influence (due to fewer competing climate modes), such as north-eastern Australia.

Therefore this paper aims to establish if there is evidence of regime shifts in the annual maxima rainfall timeseries (1-hour to 7-days) across Australia by testing for shifts (regardless of direction or timing) in the long term sub-daily and daily data. Further, the implications on IFD estimation are explored, along with the potential influence of the IPO on extreme rainfall and resulting IFD. Recommendations are then provided as to how these insights may be incorporated in future revisions of AR&R.

2. Data and methods

2.1 Data

2.1.1 Rainfall data

Sub-daily and daily rainfall data for Australia were obtained from the BoM. Sub-daily data records from continuously recording (i.e. pluviograph) rainfall stations in Australia tend to be relatively short, hindering the ability to conduct trend and attribution studies. In this study pluviograph rainfall stations were chosen with data spanning at least 40 years and at least 90% complete, resulting in 66 stations (see Figure 1a). In order to address the concerns raised
in the Introduction about short term data analysis (note that according to Raiford et al. (2007) ARI should not be extrapolated from more than twice the record length), three long-term data sets, highlighted in Figure 1a, were chosen for further analysis that contained data from at least 1913 onwards (Brisbane Aero, Sydney (Observatory Hill) and Melbourne Regional Office).

Daily rainfall stations with data spanning the period 1900 to 2009 were selected in order to capture as much temporal variability as possible (see Figure 1b). These stations were filtered according to the amount of data missing in order to identify the highest quality stations recording rainfall during this period, resulting in 96 being considered suitable for further analysis. Due to variability in the quality and quantity of rainfall data in each State of Australia, the following selection criteria were applied:

- New South Wales, Queensland and Victoria - selected stations are at least 97% complete;
- Tasmania - selected stations are at least 90% complete; and
- South Australia, Northern Territory and Western Australia - selected stations are at least 85% complete.

2.1.2 Climate index data

The climate of Australia has experienced a number of regime shifts in climate during its history, resulting in sustained periods of above average rainfall and storminess and abnormally cool temperatures, followed by the reverse conditions (i.e. droughts and elevated bushfire risk) (e.g. Erskine and Warner 1988, Franks and Kuczera 2002, Kiem et al. 2003, Kiem and Franks 2004, Verdon et al. 2004). These shifts have tended to occur every 20 to 30
years and are associated with changes in the Interdecadal Pacific Oscillation (IPO, Power et al. 1999). The IPO represents variable epochs of warming (i.e. positive phase) and cooling (i.e. negative phase) in both hemispheres of the Pacific Ocean (Folland et al. 2002). Importantly, the IPO has been shown to influence the magnitude and frequency of flood and drought cycles across eastern Australia (Kiem et al. 2003, Kiem and Franks 2004). In New Zealand, the IPO is also associated with similar shifts in flood frequency (McKerchar and Henderson 2003). It has been noted that, following the abrupt shift in the IPO in the mid 1970s, the period, amplitude, spatial structure and temporal evolution of ENSO markedly changed (Wang and An, 2001). Historically, during negative phases of the IPO there tends to be more La Niña (wet) events and fewer El Niño (dry) events (Kiem et al. 2003, Verdon and Franks 2006), resulting in an overall ‘wet’ epoch for eastern Australia and New Zealand. While during the positive phase of the IPO there tends to be a higher frequency of El Niño events and fewer La Niña events (Kiem et al. 2003, Verdon and Franks 2006), resulting in an overall ‘dry’ epoch. In this study negative phases of the IPO were defined as 1913-1920 and 1945-1977, while positive phases included 1921-1944 and 1978 to 2010.

### 2.2 Statistical tests

A 20 year moving window was used to test for low frequency variability in the annual maxima timeseries (1-hour, 1-day and 7-day). A Mann-Whitney U test was then used to determine the statistical significance of possible regime shifts by testing if the first 10 years of data was significantly different from the second 10 years, within the 20 year window (the null-hypothesis in this case was that the data was independently distributed). If the difference in medians was found to be statistically significant (i.e. p-value < 0.05) and there was a change in sign of the median values (e.g. switch from negative to positive), a climate shift was postulated to have occurred during the 10th year of the window. The Mann-Whitney U test is a robust test that does not place implicit assumptions on the underlying distribution of...
the data (i.e. it is a distribution free test), which is particularly appropriate here due to the small number of years used in each window (Kundzewicz and Robson 2004). Note that a number of different size windows were also tested, however this did not change the results or conclusions.

A second test was also applied to identify step changes in the 1-day and 7-day annual maxima time series known as the distribution free CUSUM with resampling (note that the test was not applied to the shorter sub-daily data as longer data sets are recommended for this method). CUSUM tests whether the means in two parts of a record are different (for an unknown time of change). The second test was applied as it does not require the use of a moving window (which is a limitation of the Mann-Whitney U test described above). However the CUSUM test sequentially splits the timeseries into two portions (which are not necessarily equal), which may be a problem if more than one cycle/shift is present in the timeseries.

The existence of serial correlation (or autocorrelation) in a time series will affect the ability of tests (such as the Mann-Whitney U and CUSUM) to assess the site significance of a trend (e.g. Yu et al. 2003, Serinaldi and Kilsby 2015b). The presence of cross-correlation among sites in a network will also influence the ability of the test to evaluate the field significance of trends over the network (e.g. Yu et al. 2003, Douglas et al. 2000, Guerreiro et al. 2013). Therefore, prior to applying the change point analysis as described above, the Durbin-Watson (DW) statistic was used to test for autocorrelation in the annual maxima timeseries (Durbin and Watson (1950, 1951)). In this case the null hypothesis is that the residuals from an ordinary least-squares regression are not autocorrelated against the alternative that the residuals follow an AR1 process. All DW statistic values were found to be greater than the 1.562 (the upper bound for 1% significance and a sample size of ~100) providing no evidence to reject the null hypothesis. Therefore,
any regime shifts detected using the change point methods above are not likely to be artefacts resulting from hidden persistence.

The potential issue of cross-correlation was also investigated. It was found that less than 9% of all possible pairings of rainfall data sets display a significant (yet weak) correlation at the 5% level \( r > 0.2 \), significance based on \( n=100 \). Only eight pairings (out of 4465) were correlated at 0.5 or higher. It was also found that stations located more than 500km apart were unlikely to be correlated and that the strength of the correlation reduced as distance increased between the pairs. This is not surprising given annual maximum rainfall events are due to synoptic scale processes. Therefore observations relating to spatial consistency of regime shifts are unlikely to be due to spatial correlation between sites.

2.1 IFD Calculation

The standard process for obtaining IFD information for a location is to refer to the six master charts of rainfall intensity for various durations and ARIs covering all of Australia in Volume 2 of AR&R 2001. Alternatively, IFD curves can be obtained for any location in Australia via the BoM website (both the AR&R 1987 and revised IFDs are available). This information is based on regionalised estimates of IFDs that are spatially and temporally consistent. However, this approach cannot be adopted when using the instrumental rainfall data required for the analysis presented in this study. As such, the IFD information generated for this project follows the methodology on which the IFDs were based for AR&R 1999 which utilises point source data with no regionalisation. It should be noted that it is not the purpose of this paper to compare different methods of generating IFDs, rather, one method has been adopted in order to provide a comparative assessment of the impact of non stationarity on IFD estimation. The AR&R 1999 procedure used to generate IFDs from raw rainfall data (i.e. point based estimates) is summarised as follows:
A log-Pearson III distribution was fitted to the annual maxima timeseries using the method of moments (for annual maxima series of 30 minutes to 72 hours duration). This is the standard distribution that has historically been adopted for generating IFDs in Australia; however other distributions have recently been tested as part of the revision of AR&R. To test if this distribution is suitable for the region being studied, the goodness of fit for the log-Pearson III was tested using a Kolmogorov Smirnov (KS) test. Here the null hypothesis is that the data fits the Log-Pearson III distribution (the alternate is that the data does not follow the Log Pearson III distribution). All p-values were greater than 0.05 (average p-value was 0.75), for all series (30 min to 72 hr durations at Brisbane, Sydney and Melbourne), therefore we accept the null hypothesis at the 5% significance level.

The coefficient of skewness was determined for each duration (30 minutes to 72 hours);

The coefficient of skewness was then used to obtain a frequency factor, $K_Y$, for use with Log-Pearson III Distribution. $K_Y$ was obtained from Table 2.2 (positive skew coefficients) and Table 2.3 (negative skew coefficients) in AR&R 1999 Book 4;

Rainfall intensities for a range of ARI were calculated using the following formula:

$$\log RI_Y = M + K_Y S$$  \hspace{1cm} (1)

Where:  
$RI_Y$ = rainfall intensity having an ARI of 1 in $Y$

$M$ = mean of the logarithms of the annual maxima rainfalls

$S$ = Standard deviation of the logarithms of the annual maxima rainfalls

$K_Y$ = frequency factor for the required ARI of 1 in $Y$

ARIs of 2 years to 10 years were adjusted to partial-duration series estimates. In AR&R 1999, the following correction factors were used (note: for ARI greater
than 10 years, no corrected factor is required): 2 year ARI – 1.13, 5 year ARI – 1.04, 10 year ARI – 1.0.

It should be noted that this approach is likely to result in different estimates of IFDs than those obtained from the standard maps provided by AR&R 1999 or the revised IFD estimates released in 2013. Here we are using point based rainfall data, whereas AR&R 1999 have derived regionalised estimates based on multiple rainfall stations with varying lengths of data, varying resolution (daily and pluviograph) and varying quality of records. It is recognised that analysis of rainfall data from single stations is often unreliable, is not temporally or spatially consistent and should generally not be used for design purposes. However, the use of point based rainfall data satisfies the specific aims of this study (which is a comparative analysis) and is therefore considered appropriate.

3. Results

3.1 Test for regime shifts in the annual maxima rainfall timeseries

Significant step changes identified in the extreme rainfall timeseries are shown in Figure 2. Of the 66 sub-daily rainfall stations tested, 40 (61%) displayed at least one step change in the 1-hour annual maxima timeseries (Figure 2a), with some stations exhibiting multiple shifts. Of the 96 daily rainfall stations tested, 86 displayed at least one step change in the 1-day annual maxima timeseries (Figure 2b), while 92 exhibited at least one shift in the 7-day annual maxima timeseries (Figure 2c), and some stations exhibited multiple shifts. Figure 2 collectively shows that observed step changes (or regime shifts) in annual maxima rainfall are not confined to any one particular region of Australia, with most stations analysed exhibiting at least one statistically significant shift.

****Figure 2 about here****
As shown in Figure 3, the CUSUM test yielded fewer stations with statistically significant step change in the annual maxima timeseries (only 18 stations out of 96) and many of these were clustered along the coastal fringe of eastern Australia (note that, although the total number of stations displaying significant change points was the same for both the 1-day and 7-day annual maxima, in some cases the location of the stations differed between the two). However, as stated previously a limitation to this method is that only one significant change can be detected using the CUSUM test (given that the data is sequentially split into two portions during testing). This can be a problem if more than one step change or cycle in the data is present (see example timeseries in Figure 4). Therefore, while the number of stations displaying a step change is reduced using the alternative method, the results do in fact support the theory that regime shift(s) in the annual timeseries are present for some stations at different durations.

The temporal consistency of step changes in the annual maxima timeseries was further investigated (Figure 5a) and it was found that the timing of observed shifts were not necessarily consistent across Australia. However, for some regions (e.g. the east coast of Australia) periods such as the 1940s (Figure 5b) and to a lesser degree 1970’s (Figure 5c) display a higher degree of spatial consistency. Instability and storminess can result during periods when a number of climate driving mechanisms interact (e.g. El Niño/Southern Oscillation, Indian Ocean Dipole and the Southern Annular Mode) to influence the occurrence of regional weather systems such as east
coast lows and cut off lows (Pook et al. 2006, Verdon-Kidd and Kiem 2009). However, the
large-scale climate phenomena impact various regions of Australia at different times of the
year and to varying degrees, therefore it is not surprising that the timing of shifts in the
annual maxima timeseries varies spatially and temporally. This highlights the limitations of
trying to assess and attribute variability in annual maxima rainfall based on a single climate
driver (e.g. ENSO) or attempting to address the issue of climate trends for the whole of
Australia using one simple approach or model.

3.2 Effect of non-stationarity on IFD estimation

Section 3.1 provided evidence of non-stationarity in the annual maxima timeseries for a range
of durations. This non-stationarity may ultimately influence the IFD estimation depending on
the length of data, and the time period it comes from, and therefore the underlying climatic
state (or combination of states). Current IFD estimates for Australia (both the 1987 and 2013
versions) are based on data as short as 30 years for the daily-read stations and 8 years for the
sub-daily data. Therefore IFD estimates based on relatively short-term data sets may under-
or over-estimate rainfall intensities, depending on where the data series fits within the long
term context (i.e. before or after a shift in annual maxima).

For many east coast stations a shift in 1-day annual maxima (along with the 7-day) occurred
around the 1940s - 1950s and again in the 1970s. This timing also corresponds to well-known
periods of change in the IPO (see Section 2.1.2 for a description of the IPO and its
influences). Therefore, to further explore the issue of regime shifts, breakpoints in the IPO
were used to stratify the annual maxima rainfall timeseries into IPO positive and negative
epochs for the three long sub-daily data sets described in Section 2.1.1 (i.e. Brisbane, Sydney
and Melbourne, see Figure 1a for location). The reason for selection of these stations was
twofold. Firstly, for all three stations, a shift in the annual maxima timeseries (for 1-day and
7-day) was observed during the 1940s and again in the 1970s, and secondly the stations contain long records of pluviograph data (the shortest being from 1913 onwards). Figure 6a shows the modulating effect of the IPO on total annual rainfall for the three east coast stations. Annual maxima at the three east coast stations during the two IPO epochs are also shown in Figure 6 (b-d) for event durations of 30 minutes to 72 hours (durations that are critical for flood design applications). A two-sample Kolmogorov-Smirnov (KS) test was applied to determine if the observed differences between the IPO positive and negative rainfall distributions are statistically significant. Here the null hypothesis is that the two samples are drawn from the same distribution.

**Figure 6 about here**

It is evident from Figure 6a that the effect of the IPO on annual rainfall totals (as measured by the largest difference between the two rainfall distributions associated with each climate phase and the results of the KS test) is greatest for Sydney. Although there does appear to be some impact in Brisbane, the result was not statistically significant according to the KS test. Melbourne does not appear to be greatly influenced by the IPO in terms of annual rainfall variability. This is due to the fact that the southern regions of Australia are affected by other climate modes than those arising from the Pacific (i.e. the Southern Annular Mode and the Indian Ocean Dipole (e.g. Kiem and Verdon-Kidd 2010, Gallant et al, 2012)). Regions such as Brisbane and Sydney tend to be dominated by Pacific Ocean influences (e.g. Verdon et al. 2004). Figure 6b shows annual maxima rainfall tends to be higher during IPO negative on average for durations 6 hours and longer at Brisbane (though not statistically significant according to the KS test), while Figure 6c confirms the same to be true for Sydney for durations 2 hours and longer (statistically significant at 95%). However, for Sydney, the outliers (represented by circles) tend to be larger during IPO positive, indicating that the less frequent events might be more intense during this phase.
Irrespective of the fact that the annual rainfall totals for Melbourne do not show any significant difference between the two phases of the IPO, there does appear to be a consistent relationship between IPO and the sub-daily and daily statistics (Figure 6d), whereby the median of the IPO positive distribution is higher across all durations, however IPO negative is associated with less frequent but more extreme events. (although results were not statistically significant based on the KS test). For events 24 hours and longer, the IPO negative distribution also shows a much higher degree of variability than smaller durations, with the ‘box and whiskers’ extending beyond the IPO positive counterpart for these longer durations. This suggests that while IPO might not be as dominant in southeastern Australia as it is further to the north it still has some influence that needs to be better understood.

Based on the analysis presented in Figure 6 and the results of the KS test, the Sydney record was chosen to further investigate the effects of regime shifts on IFD estimation. IFD information was generated for the Sydney record using rainfall data from the two IPO phases and the methodology outlined in Section 2.1 for durations 6 minutes through to 72 hours and ARI 2 years to 200 years. In order to test the robustness of the point estimates of rainfall return levels and estimate the uncertainty in their calculation, a simple bootstrap procedure was carried out. Firstly the IPO positive and IPO negative rainfall timeseries were resampled with replacement to obtain two new B-samples. Then for each B-sample the log-Pearson III distribution was fitted and the rainfall intensities calculated for the various return intervals. The difference between the rainfall intensities (of the two B-samples) was then calculated. This procedure was repeated 100 times to build the empirical distribution of the differences (which represents the effects of sampling and parameter estimation uncertainties under the hypothesis of the existence of two different regimes).

Figure 7 shows the difference in rainfall intensity between IPO positive and IPO negative estimates, along with the 95% confidence intervals (CIs) derived using the procedure above.
4. Discussion and conclusions

An analysis of regime shifts in the annual maxima timeseries (1-hour, 1-day and 7-day) has been carried out using a set of high quality rainfall stations across Australia. It was found that the annual maxima timeseries does indeed exhibit statistically significant step changes/shifts for the majority of stations for various durations. Further it was demonstrated using three long term sub-daily rainfall stations along the east coast that this impacts upon the resulting IFD estimation. The potential for Pacific Influences (i.e. the IPO) to influence the resulting IFD estimation was explored in order to demonstrate this issue. The authors acknowledge that the IPO is unlikely to be the only driver of variability in the annual maxima timeseries across Australia, and it is recommended that future research should aim to identify other potential drivers of this variability.

Figure 7 demonstrates clear differences in the resulting rainfall intensities for Sydney estimated for each duration and ARI using the two regimes (i.e. rainfall data from either IPO negative or IPO positive). The difference in rainfall intensity estimated is as great as 65% in some cases. In all cases, the magnitude of the difference in rainfall intensity estimated using the different data regimes is greater for less frequent events (e.g. 50-year, 100-year, 200-year ARIs), highlighting that uncertainty is greatest with less frequent events. The rainfall intensity is greater in IPO positive for the very short duration events (6 minutes) at all return intervals and for 30min duration events for return intervals of 10 years or more. Similarly, for the 24 and 72 hour duration events rainfall intensity in the positive IPO phase is higher for return intervals of 5 years or more. For 2 hour and 6 hour events, the negative phase results in higher intensity events for more frequent return levels (20 years or less) but lower intensities for less frequent events (50 years or more).
These findings highlight the fact that in some instances the IFD estimates currently being used are likely to be either under- or over-estimated at any one time depending on the length of data, and climatic state, from which they were derived. This is a particular concern given that current regionalised IFD information is based on data of varying length (as short as 8 year in the case of sub-daily data) spanning different time periods. An over estimation of rainfall intensity for a given duration could impact on construction costs, while the risks of underestimating rainfall intensities could result in failure of design criteria. That is, the risk is dependent on the application and length of time over which the risk is assessed.

Further revisions of AR&R are currently underway to include an assessment of the potential impacts of climate change on IFD estimates. However, there are many uncertainties associated with climate change projections, particularly when extracting information on timescales shorter than a season and particularly for hydrological extremes (e.g. Blöschl and Montanari 2010, Kiem and Verdon-Kidd 2011, Koutsoyiannis et al. 2008, 2009, Montanari et al. 2010, Randall et al. 2007, Stainforth et al. 2007, Stephens et al. 2012, Verdon-Kidd and Kiem 2010). Therefore, assessing future changes in extreme events that occur over short durations (e.g. minutes to days) is inherently difficult. Furthermore, climate projections are presented in terms of a percent change from a particular baseline. However, the baseline is often inconsistent and ill-defined leading to very different estimates of risk depending on the time over which the baseline is calculated (as has been demonstrated in this paper).

Importantly, for regions where large-scale climate drivers operate on a multi-year to multi-decadal timescales and are known to influence extreme rainfall events, we can use this information to determine if the climate statistics on which the IFD are based are likely to be biased or missing crucial information.

It is recommended that regime shifts in annual maxima rainfall be considered and appropriately treated in any further updates of AR&R. One way to do this may be to only...
utilise data sets of similar length ensuring that they span a sufficient number of years in order to capture data from epochs of both high or low annual maxima (to remove bias towards one climatic phase or another). However, it is acknowledged that this would potentially result in discarding a large amount of data. Alternatively, a separate set of IFDs could be developed for use in high risk modelling for engineers who need to account for the ‘worst case’ (in a similar manner to climate change allowances). This second set of IFD could be developed based on the periods of elevated annual maxima alone (for those stations with clearly defined epochs of annual maxima) such that if we were to enter such an epoch, designs based on these estimates would be robust for the duration of such a period. Salas and Obeysekera (2014) provide similar recommendations to deal with changing exceedence probabilities over time.

This would have to be assessed and calculated on a region by region basis given that Australia is a country associated with high spatial and temporal rainfall variability caused by numerous large-scale climate drivers and regional weather phenomena. Finally, any revised estimates of annual maxima should be compared in terms of uncertainty bounds (e.g. following Koutsoyiannis (2006)). Uncertainty analysis, which takes into account both the data availability and variability within the observation period would provide relevant information to practitioners about the reliability of IFD estimates.

This study has highlighted the existence of regime shifts in annual maxima rainfall data in Australia. The driving mechanisms of these regime shifts are likely to vary from location to location and decade to decade. However, these shifts are typical of many natural phenomena and can be described by processes characterized by long range dependence (or regime-switching processes) and captured by hidden Markov models (or similar), resulting in a mixture of distributions that alternate stochastically according to the transition probability from one regime to the next (e.g. Serinaldi and Kilsby, 2015a). While the strategy of defining IFDs for two (or more) different regimes (e.g. Serinaldi and Kilsby (2015a)) currently only...
partially solves the problem, as we often do not know the beginning or the end of a specific regime (be it rainfall or climate driver), recent work has focused on optimizing designs and planning strategies based on the range of what is plausible rather than a reliance on knowing the current and future climate state (e.g. Mortazavi-Naeini et al., 2015). At the same time, work is also underway on seamless prediction at a range of timescales and if/when this eventuates the results discussed here become even more important/useful. Nevertheless, the immediate usefulness of the insights presented here occurs when first establishing the IFD, as an approach similar to that employed here can be used to determine if the underlying data are biased to a mostly wet or mostly dry regime (or a mix of both) which then provides an indication as to whether the IFD is likely to be an over- or underestimate of the true risk. Importantly, this issue needs to be considered and accounted for when attempting to estimate IFD design rainfalls and prior to quantifying how those IFD estimates might change in both the near and long-term future.

While the analysis presented here has been conducted using rainfall data from Australia alone, the recommendations provided are likely to be applicable for other regions of the world where IFD information is based on short term records and particularly for locations with a highly variable climate.
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Figure 1 a) Reference stations for sub-daily stations, b) Reference stations for daily rainfall.

Note the three long term sub-daily stations used in the IFD analysis are also labelled.
Figure 2 Stations (in red) with at least one statistically significant step change in the a) 1-hour, b) 1-day, c) 7-day annual maximum rainfall (using the Mann-Whitney U test)
Figure 3 Stations (in red) with at least one statistically significant step change in a) the 1-day and b) 7-day annual maximum rainfall (using the CUSUM test with resampling)
Figure 4 Example of inadequate identification of non-stationarity using CUSUM test (red line highlights three distinct epochs of high/low rainfall, while green line demonstrates effect of splitting the data into two sections for CUSUM test)
Figure 5 a) number of stations each decade displaying evidence of a step change in 1-day annual max, b) Stations (in red) with at least one statistically significant step change in the 1-day annual max during 1940-1950 (using the Mann-Whitney U test), c) Stations (in red) with at least one statistically significant step change in the 1-day annual max during 1970-1980 (using the Mann-Whitney U test)
Figure 6 Relationship between IPO and a) total annual rainfall, and annual maximum rainfall at various durations for b) Brisbane, c) Sydney and d) Melbourne.
Figure 7 Difference in rainfall intensity for each duration and ARI. Positive (negative) values represent an increase (decrease) in rainfall intensity during IPO positive compared to IPO negative.