



1 **How streamflow has changed across Australia since**
2 **1950's: evidence from the network of Hydrologic Reference**
3 **Stations**

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5 **Sophie Xiaoyong Zhang¹, G.E. Amirthanathan¹, Mohammed Bari², Richard**
6 **Laugesen³, Daehyok Shin¹, David Kent¹, Andrew MacDonald¹, Margot Turner¹,**
7 **Narendra Kumar Tuteja³**

8 [1]{Environment and Research Division, Bureau of Meteorology, Melbourne, Australia}

9 [2]{Bureau of Meteorology, Perth, Australia}

10 [3]{Bureau of Meteorology, Canberra, Australia}

11 Correspondence to: S. X. Zhang (sophie.zhang@bom.gov.au)

12

13 **Abstract**

14 Streamflow variability and trends in Australia were investigated for 222 high quality stream
15 gauging stations having 30 years or more continuous unregulated streamflow records. Trend
16 analysis identified seasonal, inter-annual and decadal variability, long-term monotonic trends,
17 and step changes in streamflow. Trends were determined for annual total flow, baseflow,
18 seasonal flows, daily maximum flow, and three quantiles of daily flow. A distinct pattern of
19 spatial and temporal variation in streamflow was evident across different hydroclimatic
20 regions in Australia. Most of the stations in south-eastern Australia spread across New South
21 Wales and Victoria showed a significant decreasing trend in annual streamflow, while
22 increasing trends were observed in the Northern Territory and the north-west of Western
23 Australia. No trend was observed for stations in the central region of Australia. The findings
24 from step change analysis demonstrated evidence of changes in hydrologic responses
25 consistent with observed changes in climate over the past decades. For example, in the
26 Murray-Darling Basin 51 out of 75 stations were identified with step changes of significant
27 reduction in annual streamflow during the middle to late 1990s, when relatively dry years
28 were recorded across the area. Overall, the Hydrologic Reference Stations (HRS) serve as
29 'living gauges' for streamflow monitoring and changes in long-term water availability
30 inferred from observed datasets. A wealth of freely downloadable hydrologic data is provided



31 at the HRS web portal including annual, seasonal, monthly and daily streamflow data, as well
32 as trend analysis products, and relevant site information.

33

34 **Keywords:** Hydrologic Reference Stations, streamflow variability, trends, step change,
35 climate change, unregulated catchments, Australia

36

37 **1 Introduction**

38 Assessing changes and trends in streamflow observations can provide vital information for
39 sustainable water resource management. The influence of diverse environmental factors and
40 anthropogenic changes on hydrological behaviour makes the investigation into streamflow
41 changes a challenging task. Trend detection is further complicated from intra-annual, inter-
42 annual, decadal and inter-decadal variability in streamflow as well as from various
43 influencing factors that can hardly be analysed separately (WWAP, 2012; Hennessy et al.,
44 2007).

45 Extensive studies have been undertaken in different parts of the world to analyse long-term
46 hydrologic trends, and to investigate the possible effect of long-term climate variability on
47 hydrologic response (Stahl et al., 2010; Birsan et al., 2005; Lins and Slack, 2005; Milly et al.,
48 2005; Burn and Elnur, 2002). Previous works on streamflow trends draw largely on national
49 and continental analyses, especially for Europe and North America. Studies of streamflow
50 variability include analysing trends across Europe (Stahl et al., 2010; Stahl et al., 2012), and
51 at the national level. For example, Bormann et al. (2011) and Petrow and Merz (2009)
52 analysed trends under flooding conditions on German rivers. Extensive literatures on
53 hydrological trend studies have been reported for the UK: Hannaford and Buys (2012)
54 demonstrated variability in seasonal flow regimes; Hannaford and Marsh (2006, 2008)
55 analysed flow indicators at an annual resolution, and other studies focused on particular
56 regions (Biggs and Atkinson, 2011; MacDonald et al., 2010; Dixon et al., 2006; Jones et al.,
57 2006). A wide range of research on streamflow trends has been published in the USA (Kumar
58 et al., 2009; Novotny and Stefan, 2007; McCabe and Wolock, 2002) and Canada (Bawden et
59 al., 2014; Monk et al., 2011; Burn and Hag Elnur, 2002).

60 Few studies have been published for Australia to-date partly due to limited data records,
61 researches and documentation covering all flow regimes. Rivers in some regions have



62 received close attention only recently. Australia is the driest inhabited continent with an
63 average annual precipitation of 450 mm and the lowest river flow compared with other
64 continents (Poff et al., 2006). Water is relatively scarce and is therefore a valuable resource
65 across the country. Australian streams are characterized by low runoff, high inter-annual flow
66 variability, and large magnitudes of variations between the maximum and minimum flows
67 (Puckridge et al., 1998; Finlayson and McMahon, 1988). The wide variety of unique
68 topographic features combined with variable climates and frequency in weather extremes
69 result in diverse flow regimes. The recent rise in average temperature (Cleugh et al., 2011)
70 and the risk of future climate variability have added new dimensions to the challenges already
71 facing communities. Climate variability and its impact on the hydrologic cycle have
72 necessitated a growing need in Australia to seek evidence of any emerging trends in river
73 flows.

74 Chiew and McMahon (1993) examined the annual streamflow series of 30 unregulated
75 Australian rivers to detect trends or changes in the means. They found that identified changes
76 in the tested dataset were directly related to the inter-annual variability rather than changes in
77 climate. The analysis of trends in Australian flood data by Ishak (2010) indicated that about
78 30% of the selected 491 stations show trends in annual maximum flood series data, with a
79 downward trend in the southern part of Australia and an upward trend in the northern part.
80 Several other studies investigated trends of selected streamflow statistics in a particular
81 region, e.g. southwest Australia (Petroni et al., 2010; Durrant and Byleveld, 2009), southeast
82 Australia or Victoria (Tran and Ng, 2009; Stewardson and Chiew, 2009). All these studies
83 addressed the trend analysis of Australian rivers with a limited spatial or temporal coverage of
84 flow data. A gap in the research remains mainly due to constraints in access to a dataset of
85 catchments large enough to represent the diversity of flow regimes across Australia. Such a
86 dataset would enable a comprehensive and systematic appraisal of changes and trends in
87 observed river flow records.

88 The Australian national network of Hydrologic Reference Stations (HRS) was developed by
89 the Bureau of Meteorology to address this major gap and to provide comprehensive analysis
90 of long-term trends in water availability across the country (Turner et al., 2012; Zhang et al.,
91 2014). The HRS website is a one-stop portal to access high-quality streamflow information
92 for 222 well-maintained river gauges in near-natural catchments. An intention is that the



93 stations will serve as 'living gauges' that record and detect changes in hydrologic responses to
94 long-term climate variability and other factors.

95 This paper presents a statistical analysis to detect changes or emerging trends across a range
96 of flow indicators, based on the daily flow data of 222 sites from the HRS network. The
97 objective of this study is to provide a nationwide assessment of the long-term trends in
98 observed streamflow data. Evaluation of past streamflow records and documenting recent
99 trends will be of benefit in anticipating potential changes in water availability and flood risks.
100 It is hoped that the findings from trend analysis presented in this paper will inform decision
101 makers on long-term water availability across different hydroclimatic regions, and be used for
102 water security planning within a risk assessment framework.

103

104 **2 Site selection, data and methods**

105 **2.1 Hydrologic Reference Stations and data**

106 The 222 Hydrologic Reference Stations (HRS) were selected from a preliminary list of
107 potential streamflow stations across Australia according to the HRS selection guideline (SKM
108 2010). These guidelines specified four criteria for identifying the high quality reference
109 stations, namely unregulated catchments with minimal land use change, a long period of
110 record (greater than 30 years) of high quality streamflow observations, spatial
111 representativeness of all hydro-climate regions, and the importance of site as assessed by
112 stakeholders. The station selection guidelines were then applied in four phases
113 (www.bom.gov.au/water/hrs/guidelines.shtml). The HRS network will be reviewed and
114 updated every two years to ensure that the high quality of the streamflow reference stations is
115 maintained.

116 Two features were considered in order to define the hydroclimatic regions in HRS: climatic
117 zones and Australia's drainage divisions. The climatic zones were defined according to
118 climate classification of Australia based on a modified Koeppen classification system (Stern
119 et al., 2000). Australia has a wide range of climate zones, from the tropical regions of the
120 north, through the arid expanses of the interior, to the temperate regions of the south (ABS
121 2012). The Australian Hydrological Geospatial Fabric (Geofabric) Surface Catchments (BOM
122 2012) were used to delineate 12 topographically defined drainage divisions approximating the
123 drainage basins from the Geoscience Australia (2004) definition. The selection of HRS



124 stations aimed to maximise the geographical extent of the available records. As shown in
125 Figure 1, the final set of 222 hydrologic reference stations cover all climatic zones,
126 jurisdictions and most drainage divisions. Since most Australian rivers are located near the
127 coast, there is a high density of stations along the coast and sparsely distributed stations across
128 inland areas. One third of the HRS sites are in temperate climate zone, and the majority of the
129 rest are either in Tropical or Subtropical regions; only a few are located in other climate
130 zones. The distribution of Hydrologic Reference Stations across multiple hydroclimatic
131 regions provides data for a comprehensive investigation of long-term streamflow variability
132 across Australia.

133 The primary data used in this study were daily streamflow series of 222 gauging stations from
134 the HRS network. Table 1 lists the twelve drainage divisions and the number of stations in
135 each division. One third of the HRS stations are located within the Murray-Darling basin, half
136 of the rest are distributed along eastern coasts. This is the best compiled long-term quality
137 controlled data for Australia and the trends derived from this dataset constitute the first such
138 statement on long-term water availability across Australia.

139 The earliest record included in the data set is from 1950. Data prior to this has been excluded
140 due to the common existence of large gaps in the pre-1950 period. All stations included in the
141 HRS had a target of 5% or less missing data to meet the completeness criteria for high quality
142 streamflow records. Some stations were included with more than 5% missing data where they
143 excelled in other criteria such as stakeholder importance or spatial coverage. The periods of
144 data gaps were filled using a lumped rainfall-runoff model GR4J (Perrin et al. 2003), which
145 was found to perform well at most sites. The model was calibrated and forced with catchment
146 average rainfall and potential evapotranspiration from the Australian Water Availability
147 Project (AWAP) (Raupach et al., 2009).

148 The study examined sites with varying lengths of record depending on the data availability.
149 The daily flow data were aggregated into annual series based on a water year calculation. The
150 start month of the water year was defined as the month with the lowest monthly flow across
151 the available data period. In order to ensure the statistical validity of the trend analysis, all
152 stations had minimum 30 years of record, with an average time-series length of 45 years. The
153 longest record length was 62 years, 25% of the stations have 50 or more years of record.
154 Catchment sizes ranged from 4.5 to 232,846 km² with a mean size of 3108 km². The majority



155 (82%) of the stations had an upstream drainage area less than 1000 km², and only three
156 stations had a drainage area larger than 50,000 km².

157 The data and the long term series gathered in this study are the best compiled and quality
158 assured data for HRS catchments. The analysis and trends derived from the HRS datasets
159 constitute the first statement on long-term water availability across Australia.

160 **2.2 Streamflow variables for trend analysis**

161 Long-term climate variability can be reflected through trends in streamflow variables. To
162 understand the importance of the components of the hydrologic regimes and their potential
163 link to long-term climate variability, ten streamflow variables were chosen for statistical and
164 trend analysis. Two variables related to fluctuation of annual flows were annual total flow
165 (Q_T) and annual baseflow (Q_{BF}). Baseflow was separated from daily total streamflow using a
166 digital filter based on theory developed by Lyne and Hollick (1979) and applied by Nathan
167 and McMahon (1990).

168 Daily streamflow data were analysed to form a group of indicators of daily flow trends. They
169 were daily maximum flow of each year (Q_{Max}), the 90th percentile (non-exceedance
170 probability) daily flow of each year (Q_{90}), the 50th percentile daily flow of each year (Q_{50}),
171 and the 10th percentile daily flow of each year (Q_{10}). The median daily flow Q_{50} was used in
172 the study instead of daily mean flow because the flow distribution is skewed and outliers are
173 present.

174 Four seasonal total flow indicators were analysed to examine the seasonal trend patterns.
175 These variables included summer flow Q_{DJF} (December to February), autumn flow Q_{MAM}
176 (March to May), winter flow Q_{JJA} (June to August), and spring flow Q_{SON} (September to
177 November).

178 The trend analysis was applied to the ten hydrologic indicators of streamflow data at each
179 HRS station.

180 **2.3 Trend and data statistical analysis**

181 Changes in streamflow data can occur gradually or abruptly. Statistical significance testing is
182 commonly used to assess the changes in hydrological datasets (Helsel and Hirsch, 2002;
183 Monk et al., 2011; Hannaford and Buys, 2012). The Mann-Kendall (MK) trend test (Mann,



184 1945; Kendall, 1975) was adopted in this study to identify statistically significant monotonic
185 increasing or decreasing trends (Petroni et al., 2010; Zhang et al., 2010; Miller and Piechota,
186 2008). In order to ensure the assumption of independence was met for the MK test, the non-
187 parametric Median Crossing and Rank Difference tests (Kundzewicz and Robson, 2000) were
188 applied to entire datasets. When either of the randomness tests indicated that the time series
189 was not from a random process, the site was excluded from the MK trend assessment. As this
190 study attempted to examine patterns in historical streamflow records, no further adjustments
191 were made to account for the non-random structure of data.

192 The non-parametric MK trend test was used to detect the direction and significance of the
193 monotonic trend, and the trend line from a least squares regression (LSR) was used to
194 approximately represent the magnitude of the trend. The trend magnitude was standardised
195 [trend (mm/yr) / average annual flow (mm)] to make the change comparable across stations
196 by dividing the regression slope coefficient by the average annual flow over the data period.

197 All data were subject to step change analysis to detect any abrupt changes during the record
198 period. The distribution free CUSUM test (Chiew and Siriwardena, 2005) was applied to
199 identify the year of change in streamflow series. The significant difference between the
200 median of the streamflow series before and after the year of change was tested by Rank-Sum
201 method (Zhang et al., 2010; Miller and Piechota, 2008; Chiew and Siriwardena, 2005). More
202 information on the statistical tests used in this study can be found in Appendix A.

203 In addition to the trend analysis for the ten flow indicators, other statistical data analyses were
204 included to gain a broad understanding of hydrologic regimes. Aggregated monthly and
205 seasonal flow data were investigated for changes in flow patterns in different basins or
206 regions. Daily event frequency analyses were used to examine the variations in daily
207 streamflow magnitude, and daily flow duration curves were presented to examine changes in
208 daily flow among decades.

209

210 **3 Development of the HRS web portal**

211 A web portal has been developed to house the network of Hydrologic Reference Stations and
212 provide access to streamflow data, results of analysis, and associated site information. Figure
213 2 summarises the development process of the HRS network and website. Through a data
214 quality assurance process following the guidelines and stakeholder consultations, the final list



215 of 222 streamflow gauging stations was established. A suite of software tools, "the HRS
216 toolkit" was developed to undertake data aggregation, analysis, trend testing, visualisation and
217 manipulation. The toolkit is capable of automatically converting the flow variables to
218 monthly, seasonal and annual totals, and quantifying the step and/or linear changes in the
219 selected streamflow variables. The toolkit also generated and processed graphical products,
220 data, statistical summary tables and statistical metadata included in the web portal.

221 A snapshot of the HRS web portal is shown in Figure 3. The main page was designed with
222 three parts. A series of links on the top provide the project information. Below this is the
223 station selector, which facilitates searching for the site of interest by location. The third part is
224 the product selector containing the core information sections of the website. Several tabs are
225 offered for users to explore the web portal dependent on their needs and the level of
226 information they require. The daily streamflow data, graphical products, statistics and trend
227 analysis results are available for users to view and download. Information provided on the
228 HRS web portal will assist in detecting long-term streamflow variability and changes at the
229 222 sites, and therefore supports water planning and decision-making. More information can
230 be found at the website <http://www.bom.gov.au/water/hrs>.

231 This web portal provides public access to high quality data and information. It has more than
232 15,000 graphic products for display. It is carefully designed for the public to have synthesised
233 and easily understandable information on water availability trends across Australia. In order
234 to ensure currency of this web site, streamflow data are updated and reviewed every two
235 years.

236

237 **4 Result and discussion**

238 The study to detect long-term streamflow trends was performed on the 222 gauging stations
239 included in the HRS network. This section presents an overview bar-plot of the Mann-Kendall
240 test results for the selected ten hydrologic variables. Maps showing trend detection results and
241 step change analysis for the annual total flow are presented as well as a table listing the
242 stations with significant trends in annual total flow at 1% significance level ($p < 0.01$). In
243 addition, variations in trend among daily flow indicators and seasonal flows are examined.
244 Finally, regional patterns in long term trends, inter-annual and decadal variability are further
245 investigated for two feature stations.



246 **4.1 Overview**

247 A stacked bar-plot is shown in Figure 4 that stratifies the stations by the trend across each
248 streamflow variable. Overall, a consistent pattern is seen across the 10 streamflow variables –
249 the majority of stations have either no trend or a non-random time-series; of the stations with
250 a trend detected, the majority are decreasing.

251 A distinction was noted between patterns of trends in the different flow regimes. Moving
252 through the flow variables from low, to median, to high, and onto maximum, an increasing
253 number of stations were found with no trends. The overall number of stations with
254 statistically significant trends was around the same across the median, high, and maximum
255 variables but much lower for the low flow variable. Around one third of stations showed a
256 decreasing trend in spring and a quarter of stations in summer and winter. A significant
257 proportion of stations do show a decreasing trend across most variables. Summer had a large
258 number of stations with no trend and 3 stations with an increasing trend. Due to non-
259 randomness of streamflow variables, a number of stations are not amenable to trend analysis.

260 **4.2 Spatial distribution of trends in annual total streamflow**

261 Detecting the trend and non-stationarity in a hydrologic time series may help us to understand
262 the possible links between hydrological processes and global environment changes. Many
263 hydrological time series exhibit trend or non-stationarity in the mean or median. The long-
264 term gradual change in rainfall-runoff transformations could be represented by linear trend.
265 The abrupt changes in a hydrologic time series could be due to hydrologic non-stationarity.

266 **4.2.1 Linear trend**

267 Maps were generated showing the trend results for each variable across Australia. The trend
268 analysis map of annual total streamflow (Q_T) displays the direction and significance of a trend
269 (Figure 5) at different levels of significance: $p < 0.01$, $p < 0.05$ and $p < 0.1$. Although trends
270 in Q_T vary across different hydro-climatic regions of the continent, a clear spatial pattern is
271 evident from the map: about 35% of the stations showing decreasing trends are in the
272 southern part of Australia and 4% increasing trends in the northern part, while there no
273 significant trend visible in the central region of Australia. The general downward trends
274 observed in southern Australia may have been affected by the dry period in the last decade in
275 the south-eastern and south-western regions. Stations in the Murray-Darling Basin



276 demonstrated the strongest decreasing trends with 30 stations exhibiting high levels of
277 significance at $p < 0.05$.

278 A set of 22 gauging stations were identified with trends in annual total streamflow at 0.01
279 significance levels, see Table 2. All sites showed consistent direction of change using MK test
280 and LSR. None of those 22 gauges showed increasing trend. Trends in annual baseflow were
281 found to be similar to the results of annual totals when a significant trend was detected. The
282 number of stations showing significant declining trends in baseflow conditions was less than
283 it was for annual total flow. However, some time-series of annual baseflow were non-random
284 and therefore not available for further trend testing.

285 Step change analysis was applied to all sites where the time series data was random to give
286 comparable results of gradual and abrupt changes in annual total flows. Table 2 gives the
287 Rank-Sum test results and lists the year of change for the 22 stations. Details of step changes
288 across Australia will be discussed in the following section.



289 **4.2.2 Step change**

290 The Rank-Sum test was used to identify the presence of a step change in the median of two
291 periods, with the distribution free CUSUM method providing the year of change. Values were
292 reported for sites with Rank-Sum test at 0.1 significance levels or higher. Figure 6 shows the
293 results of step change analysis, where colours indicate the year of change appearing in various
294 decades, and upward arrows represent increased median values after the year of change and
295 vice versa.

296 The step change map reveals a definite spatial pattern in the location of stations that exhibited
297 a significant step change. As expected, the direction and significance of step-changes is
298 consistent with the Mann-Kendall results for most stations. The identified years of step
299 changes appear to show spatial groupings at different divisions. The majority of stations in
300 southeast Australia were characterised with step changes in mid-1990s, when the millennium
301 drought (BOM and CSIRO, 2014) started to dominate the weather in this region. Five stations
302 in south-west West Australia had a key feature of 1975 step change, which might be partly
303 due to the observed rainfall decline since the mid-1970s. It was also noted that most stations
304 located on the south east coast of Queensland showed a significant step change in the 1980s.

305 The results from step change analyse imply that changes in streamflow and the consequent
306 hydrologic response are driven by changes in climatic forcing such as rainfall over the period
307 of record. Investigating this causative relationship and quantifying the relative impacts of
308 variations in climate on streamflow predictions is left for future work.

309 **4.3 Spatial distribution of trends in daily flows and seasonal flows**

310 Trend analysis maps shown in Figure 7 decompose trends of daily flow for Q_{Max} , Q_{90} , Q_{50} and
311 Q_{10} . In general, the identified trends were spatially consistent with the trend pattern in Q_T :
312 with upward trends in the north-west and downward trends in the south-east, south-west and
313 Tasmania. The Q_{50} and Q_{10} series are notable for the number stations with non-random time-
314 series and therefore an invalid MK test result, this can be seen most dramatically in Figure 7d,
315 and is due to the higher correlation of the time-series. This daily flow trend analysis indicated
316 similar results to previous studies (Tran and Ng, 2009; Durrant and Byleveld, 2009) for the
317 respective sites and flow statistics.



318 The analysis of maximum daily flow Q_{Max} could be considered as analysis of extreme flow as
319 this series contains the maximum value for each year. The general pattern of trends in Q_{Max}
320 was in accordance with the preliminary trend analysis results in Ishak (2010), which
321 suggested that about 30% of selected stations showed trend in Q_{Max} , with downward trend in
322 the southern part of Australia and upward trends in the northern part (Figure 7a).

323 The spatial distribution of trends in the seasonally disaggregated total flow series were
324 investigated (Figure 8). The broad pattern from the analysis is a collection of downward
325 trends generally in the south and upward trends in the north across the seasonal variables;
326 summer (Q_{DJF}), autumn (Q_{MAM}), winter (Q_{JJA}), and spring (Q_{SON}). However, contrasting
327 Figure 5 and Figure 8 suggest that the trends detected in the annual total flows series are
328 predominantly a mixed result of increasing summer trends in northern Australia, and
329 decreasing winter trends for southern Australia.

330

331 5. Discussion

332 A comprehensive statistical and trend analysis in long-term streamflow data was conducted
333 for 222 unregulated river gauges from the HRS national network. Ten streamflow variables
334 were examined to detect underlying changes or trend in streamflow and to identify spatial
335 variations across Australia.

336 Commonality and differences were found from this study when compared with previous
337 streamflow trend studies across Australia. This could be expected given the different selection
338 of flow statistics, gauge location, data length, employed techniques and methodology. For
339 example, to examine the trends in south-west Western Australia (SWWA), Durrant and
340 Byleveld (2009) has investigated 29 sites in the area using post-1975 data, whilst this paper
341 considered the full record of data since 1950 and the full water year was used. Owing to the
342 different data record periods used in trend analysis, seven stations in Durrant and Byleveld
343 (2009) showed a possible increase, while in this study a homogenous spatial pattern of
344 downward trends was revealed across the SWWA. Three stations in common were examined
345 by both studies. The streamflow data of Yarragil Brook at Yarragil Formation (614044) was a
346 non-random series, which was strongly biased by the 1975 step change. When only looking at
347 the runoff of post-1975 period at this site, it revealed a very weak decreasing trend, which was
348 similar to the result of Durrant and Byleveld (2009). Carey Brook at Staircase Road (608002)
349 had similar time series data starting from the mid-1970s in both studies. A slight decreasing



350 linear trend and a 1997 step change at 0.05 significance level was identified in this study. No
351 statistically significant trend was detected in Durrant and Byleveld (2009), which could be
352 attributed to the limited record until 2008 and not considering the recent years of 2010, 2011
353 and 2012 that were relatively dry. The results were in agreement in both studies showing no
354 strong decreasing trend for the Kent River at Styx Junction (604053). At this site the 1975
355 change was not predominant.

356 The results of this study have demonstrated the main characterisation of hydrological change
357 of river flows across Australia since the 1950's. Overall, most of the downward trends in Q_T
358 appeared within or very close to the temperate climate zone, while upward trends were in the
359 tropical region. The spatial pattern of trends matched the rainfall records maps that indicated
360 rainfall deficiency in the south in the last decade comparing the historical records (Cleugh et
361 al., 2011). Similar rainfall changes were also observed all over the continent as shown in the
362 recent CSIRO sustainable yield study projects (CSIRO, 2013). Drought conditions persisted
363 in the south-east and south-west of the continent from around 1996 to 2010 might be
364 attributed to the detected change in streamflow. This could be the reason that most of the
365 gauging stations in southern Australia and southeast of Queensland showed a significant
366 decreasing trend in annual streamflow. It was also found that positive trends observed at
367 many locations in northern Australia could be related to increased rainfall in this part of
368 Australia during the last decade. Other changes such as within-year rainfall variation and
369 increase in temperature may have played a role in affecting the hydrologic cycle.

370 Whilst it is a possible explanation, it is not explicit that climate change is the cause of
371 significant trends in streamflow. There are many other factors that may affect streamflow, for
372 example, natural catchment changes, climate variability, data artefacts and other influences.
373 Site specific comparison of rainfall, PE, and temperature may help to improve the
374 understanding of the underlying causes of trends in hydrological variables. Further
375 investigation would be required to discover the potential causes of detected trends, which was
376 beyond the scope of this study.

377 Under the Water Act (2007), the Australian Bureau of Meteorology has responsibility for
378 compiling and disseminating comprehensive water information nation-wide. Hydrologic
379 Reference Stations (HRS) is an initial step to build up the national river data network. The
380 network of HRS, which the present study was based on, is the first operational website in
381 Australia as a national river flow data repository. It provides an excellent foundation for water



382 planning and research – particularly in trend detection and the possibility to link to large scale
383 atmospheric and climate variables. The information on the HRS website can be used as a test
384 bed for model development, hydrological non-stationarity assessments and many other
385 research interests.

386

387 **6. Conclusions**

388 This study investigated the streamflow variability and inferred trends in water availability for
389 222 gauging stations in Australia with long term and high quality streamflow records. The
390 results present a systematic analysis of recent hydrological changes in greater spatial and
391 temporal details than previously published for Australian rivers. Implications of the findings
392 should aid decision making for water resources management, especially when considering the
393 results in the context of climate variability.

394 The main findings of the study are:

395 • The spatial and temporal trends in observed streamflow varied across different
396 hydro-climatic regions in Australia (Figure 2). In Northern Territory and north-west of
397 Western Australia, there was an increasing trend in annual streamflow (Q_T) while
398 there was no significant trend visible in the northern region of Queensland. However,
399 in south-eastern Queensland there was a significant decreasing trend. Most of the
400 gauging stations in New South Wales, Victoria and north-west Tasmania showed a
401 significant decreasing trend in annual streamflow. In South Australia and South
402 Tasmania, most of the stations showed no significant trend in annual streamflow.

403 • The temporal trends also varied between different components of streamflow –
404 annual total, daily maximum (Q_{Max}), high, median and low flows (Q_{90} , Q_{50} , Q_{10}),
405 baseflow (Q_{BF}) and seasonal totals (Q_{JJA} , Q_{SON} , Q_{DJF} , Q_{MAM}). Out of 222 stations, only
406 7 showed an increasing trend, 90 decreasing and 98 no trend in total annual
407 streamflow. The annual daily maximum streamflow showed decreasing trends at 67
408 stations while the low flow and baseflow components showed decreasing trends at 18
409 and 73 stations respectively. Trends also varied between different seasonal totals and
410 also across different hydro-climatic regions. Most of Northern Territory and central
411 Australia showed increasing trend in summer (Q_{DJF}) flow while no stations were found
412 with increasing trend for winter flow (Q_{JJA}) anywhere in Australia.



413 • The analysis of step changes revealed definite regional patterns: stations in
414 southeast Australia were characterised with step changes in the mid-1990s, while a
415 key feature of a 1975 step change was identified for stations in south-west West
416 Australia.

417 • The web portal (<http://www.bom.gov.au/water/hrs>) displays all the graphical
418 products, tables, and statistical test results of all 222 stations. It contains a
419 comprehensive unique set of graphical products for linear trends and step change.

420 The streamflow trends evident from the statistical data analysis showed some parallels with
421 climate variability patterns that the country experienced through recent decades. Long-term
422 trends in water availability across different hydroclimatic regions of Australia reported in this
423 study are derived purely from observations unlike other studies, they are not derived from
424 models which can invariably be influenced by biases. The high quality streamflow data of
425 HRS and the results from this analysis on streamflow variability provide critical information
426 for water security planning and for prioritising water infrastructure investments across
427 Australia.

428

429 **Appendix A: Statistical tests**

430 **A1. Median Crossing Test**

431 This method tests for randomness of a time series data. It is a non-parametric test. The n time
432 series values ($X_1, X_2, X_3, \dots, X_n$) are replaced by '0' if $X_i < X_{\text{median}}$ and by '1' if $X_i > X_{\text{median}}$. If
433 the time series data come from a random process, then the count 'm', which is the number of
434 times 0 is followed by 1 or 1 is followed by 0, is approximately normally distributed with:

435 Mean: $\mu = \frac{(n-1)}{2}$

436 Standard deviation: $\sigma = \frac{(n-1)}{4}$

437 The z-statistic is therefore defined as:



438
$$z = \frac{|(m - \mu)|}{\sigma^{0.5}}.$$

439 **A2. Rank Difference Test**

440 This method also tests for randomness of a time series data. It is a non-parametric test. The n
 441 time series values ($X_1, X_2, X_3 \dots X_n$) are replaced by their relative ranks starting from the
 442 lowest to the highest ($R_1, R_2, R_3 \dots R_n$). The statistic ‘U’ is the sum of the absolute rank
 443 differences between successive ranks:

444
$$U = \sum_{i=2}^n |R_i - R_{i-1}|$$

445 For large n, U is normally distributed with:

446 Mean:
$$\mu = \frac{(n+1)(n-1)}{3}$$

447 Standard deviation:
$$\sigma = \frac{(n-2)(n+1)(4n-7)}{90}$$

448 The z-statistic* is therefore defined as:

449
$$z = \frac{|(U - \mu)|}{\sigma^{0.5}}.$$

450 **A3. Mann-Kendall Test**

451 This method tests whether there is a trend in the time series. It is a non-parametric rank-based
 452 test. The n time series values ($X_1, X_2, X_3 \dots X_n$) are replaced by their relative ranks starting
 453 from the lowest to the highest ($R_1, R_2, R_3 \dots R_n$).

454 The test statistic S is defined as:

455
$$S = \sum_{i=1}^{n-1} \left[\sum_{j=i+1}^n \text{sgn}(R_i - R_j) \right]$$

456 where $\text{sgn}(y) = 1$ for $y > 0$
 457 $\text{sgn}(y) = 0$ for $y = 0$
 458 $\text{sgn}(y) = -1$ for $y < 0$
 460 $\text{sgn}()$ is the signum function.

461 If there is a trend in the time series (i-e the null hypothesis H_0 is true), then S is
 462 approximately normally distributed with:

463 Mean:
$$\mu = 0$$

464 Standard deviation:
$$\sigma = \frac{n(n-1)(2n+5)}{18}$$

465 The z-statistic* is therefore:



$$z = \frac{|S|}{\sigma^{0.5}}$$

467 A positive value of S indicates that there is an increasing trend and vice versa.

468 **A4. Distribution Free CUSUM Test**

469 This method tests whether the means in two parts of a record are different for an unknown
470 time of change. It is a non-parametric test. Given a time series data ($X_1, X_2, X_3 \dots X_n$), the test
471 statistic V_k is defined as:

472

$$473 \quad V_k = \sum_{i=1}^k \text{sgn}(X_i - X_{\text{median}})$$

474

475 where $\text{sgn}(y) = 1$ for $y > 0$
476 $\text{sgn}(y) = 0$ for $y = 0$
477 $\text{sgn}(y) = -1$ for $y < 0$
478 X_{median} is the median value of the X_i data set.

479 The time at which ' $\max|V_k|$ ' occurs is considered as the time of change. The distribution of V_k
480 follows the Kolmogorov-Smirnov two-sample statistic ($KS = (2/n) \max|V_k|$). A negative value
481 of V_k indicates that the latter part of the record has a higher mean than the earlier part and vice
482 versa.

483 **A5. Rank-Sum Test**

484 This method tests whether the medians in two different periods are different. It is a
485 nonparametric test. The time series data is ranked to compute the test statistic. In the case of
486 ties the average of ranks are used. The statistic S is the sum of ranks of the observations in the
487 smaller group. The theoretical mean and standard deviation of S under H_0 for the entire
488 sample is given as:

$$489 \quad \text{Mean:} \quad \mu = \frac{n(N+1)}{2}$$

$$490 \quad \text{Standard deviation:} \quad \sigma = \left[\frac{nm(N+1)}{12} \right]^{0.5}$$

491 where n and m are the number of observations in the smaller and larger groups
492 respectively. The standardised form of the test statistic, Z^* is computed as:

$$493 \quad Z = (S - 0.5 - \mu) / \sigma \quad \text{if } S > \mu$$

$$494 \quad Z = 0 \quad \text{if } S = \mu$$

$$495 \quad Z = |S + 0.5 - \mu| / \sigma \quad \text{if } S < \mu$$

496 Z is approximately normally distributed.



497 **Acknowledgements**

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504

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632 Table 1: Metadata of the drainage divisions and selected Hydrologic Reference Stations

Division map code	Drainage division names	Mean annual rainfall (mm) (1976-2005)*	Mean elevation (m)	Number of HRS stations	Water year start month	Smallest catchment area (km ²)	Largest catchment area (km ²)
I	Northeast Coast	764	173	42	October	6.6	7486.7
II	Southeast Coast	599	323	44	March	4.5	4660.0
III	Tasmanian	1519	199	12	February	18.3	775.3
IV	Murray-Darling	479	260	75	March	26.3	35238.9
V	South Australia Gulf	344	269	5	February	5.3	187.4
VI	Southwest Coast	329	365	13	March	14.1	1786.0
VII	Indian Ocean	369	162	0	(No data)	(No data)	(No data)
VIII	Timor Sea	520	339	13	September	65.4	47651.5
IX	Gulf of Carpentaria	674	293	13	October	170.0	43476.2
X	Lake Eyre	429	312	5	October	434.9	232846.3
XI	North Western Plateau	456	359	0	(No data)	(No data)	(No data)
XII	South Western Plateau	321	297	0	(No data)	(No data)	(No data)

633 * Calculation was based on rainfall data from BOM climate website <http://www.bom.gov.au/>



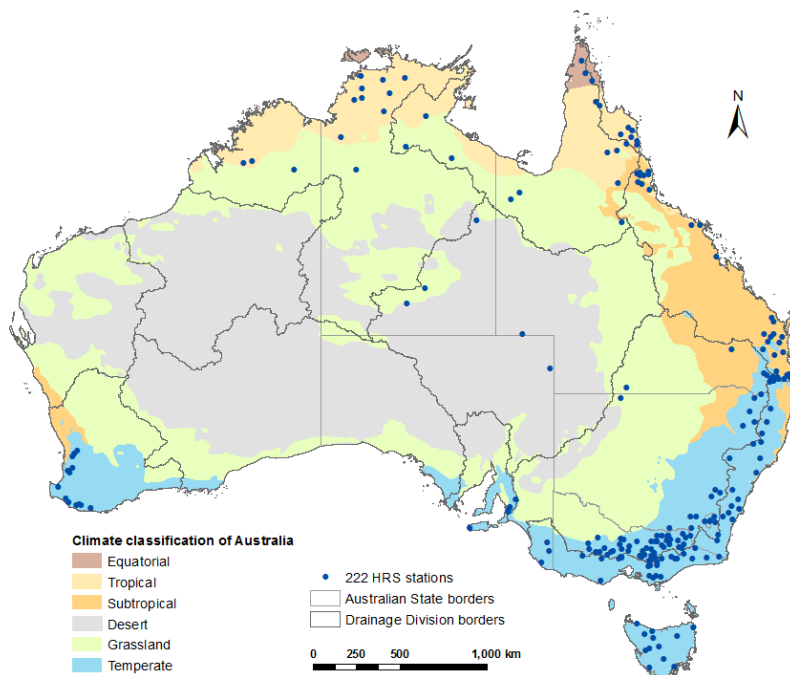
Table 2: Results of trend analysis for stations showing MK trend test at 0.01 significance level in annual total streamflow

Div	State	Basin	Station ID	Site name	Area (km ²)	Data time series		Ave. annual flow (GL/yr)	BF index	Trend		Step change	
						Start year	End year			MK	LSR	Rank	Year
II	VIC	Snowy River	222206	Buchan River at Buchan	850	1951	2014	140.0	0.45 ⁻	** ↓	-1.5%	**	1976
	VIC	Mitchell-Thomson Rivers	223202	Tambo River at Swifts Creek	899	1951	2014	77.1	0.46 ⁻	** ↓	-1.8%	**	1978
	VIC	Wentbee River	231213	Lerderg River at Sardine Creek Obrien Crossing	152	1959	2014	25.9	0.34 ⁻	** ↓	-1.9%	**	1996
	VIC	Hopkins River	236213	Mount Ernu Creek at Mena Park	448	1974	2014	13.6	0.18 ⁻	** ↓	-3.5%	**	1996
	VIC	Glenelg River	238208	Jimmy Creek at Jimmy Creek	23	1951	2014	3.4	0.47 ^{** ↓}	** ↓	-1.8%	**	1996
	SA	Millicent Coast	A2390519	Mosquito Creek at Struan	1550	1971	2014	21.7	0.25 ⁻	** ↓	-3.2%	**	1992
	SA	Millicent Coast	A2390523	Stony Creek at Wookvime Range	485	1973	2014	4.8	0.55 ^{** ↓}	** ↓	-2.9%	**	1990
III	TAS	Smithton-Burnie Coast	314213	Black River at South Forest	318	1968	2014	194.1	0.38 ⁻	** ↓	-1.0%	**	1992
IV	NSW	Upper Murray	401009	Maragle Creek at Maragle	217	1951	2014	35.9	0.47 ^{** ↓}	** ↓	-1.6%	**	1996
	VIC	Kiewa River	402217	Flaggy Creek at Myrtleford Road Bridge	26	1970	2010	4.0	0.42 ^{** ↓}	** ↓	-2.5%	**	1996
	VIC	Goulburn	405238	Mollison Creek at Pyalong	164	1972	2014	19.5	0.29 ⁻	** ↓	-3.5%	**	1996
	VIC	Goulburn	405248	Major Creek at Graytown	288	1971	2014	13.2	0.10 ^{** ↓}	** ↓	-4.2%	**	1996
	VIC	Goulburn	405251	Brankeet Creek at Ancona	122	1973	2014	14.8	0.45 ⁻	** ↓	-2.2%	**	1996
	VIC	Campaspe River	406214	Axe Creek at Longlea	237	1972	2014	13.4	0.18 ⁻	** ↓	-4.0%	**	1996
	VIC	Loddon River	407214	Creswick Creek at Chunes	300	1951	2014	24.0	0.32 ^{** ↓}	** ↓	-2.1%	**	1996
	VIC	Loddon River	407230	Joyces Creek at Strathlea	156	1973	2014	9.2	0.17 ⁻	** ↓	-3.0%	**	1996
	NSW	Lachlan	412028	Abercrombie River at Abercrombie	2631	1951	2014	277.0	0.30 ⁻	** ↓	-2.0%	**	1978
	NSW	Lachlan	412066	Abercrombie River at Hadley No.2	1630	1960	2014	169.8	0.29 ⁻	** ↓	-2.1%	**	1978
	VIC	Avon	415226	Richardson River at Carrs Plains	125	1971	2014	3.7	0.04 ^{** ↓}	** ↓	-4.3%	**	1996
	VIC	Wimmera	415237	Concongella Creek at Stawell	244	1976	2014	9.1	0.12 ^{** ↓}	** ↓	-4.6%	**	1996
VI	WA	Murray River (WA)	613002	Harvey River at Dingo Road	148	1970	2014	29.7	0.58 ⁻	** ↓	-1.8%	**	1993
	WA	Swan Coast	616065	Canning River at Glen Eagle	537	1953	2014	18.9	0.36 ^{** ↓}	** ↓	-2.7%	**	1975

* Significant at $p < 0.05$ ** Significant at $p < 0.01$

- baseflow series non-random ° baseflow no trend

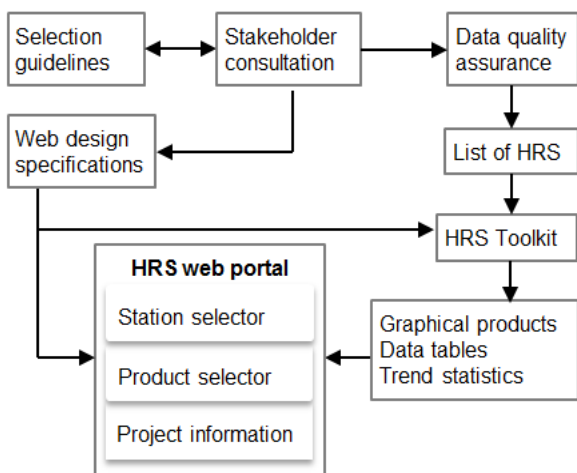
↓ decrease trend ↑ increase trend



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637 Figure 1: Location of the 222 high quality streamflow reference stations, the climatic regions
638 and Australia drainage divisions (Geofabric Surface Hydrology Catchments, Geofabric V2.1,
639 BOM 2012)

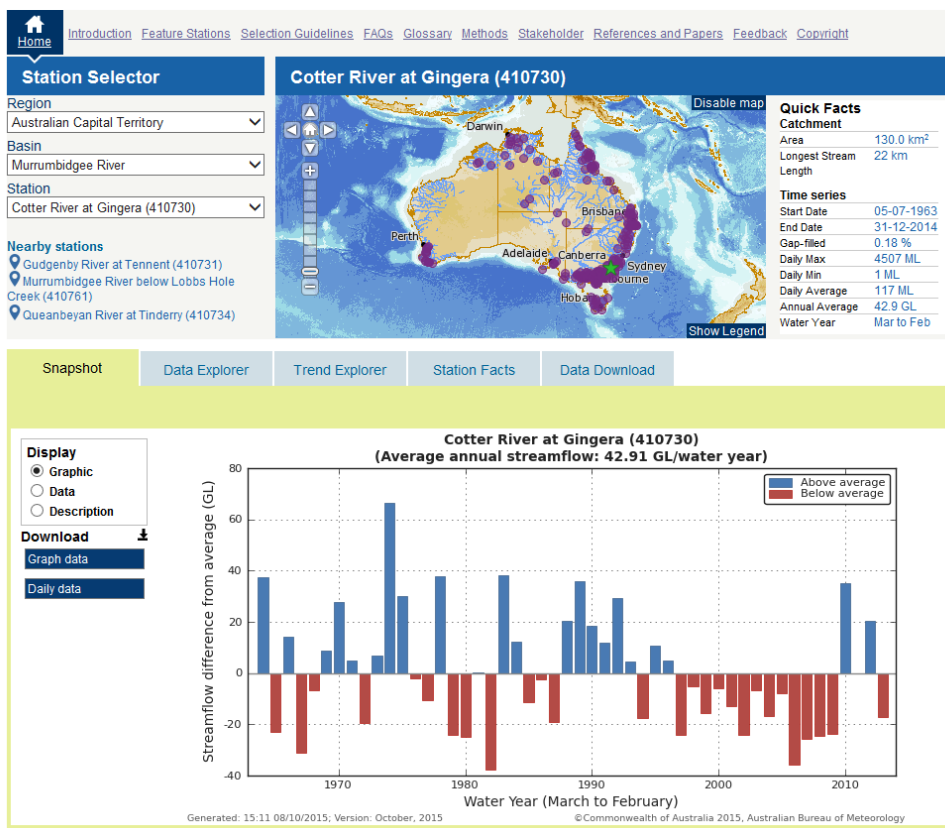


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642 Figure 2: Framework of developing Hydrologic Reference Stations

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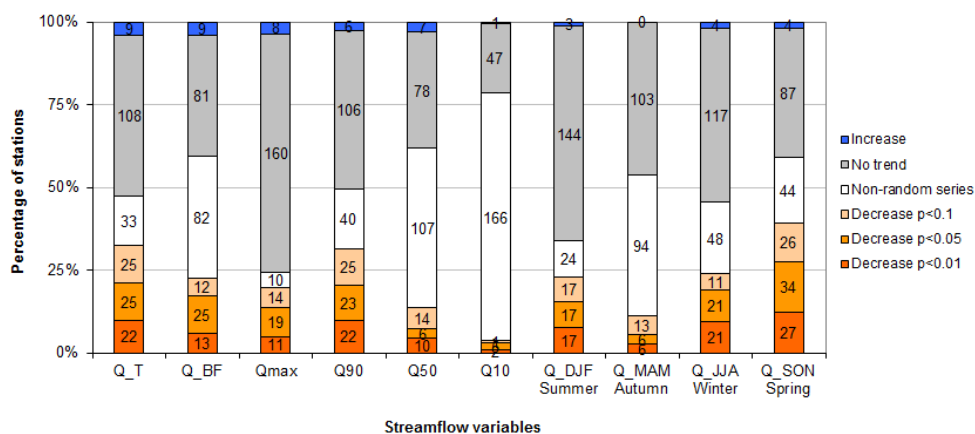


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646 Figure 3: Snapshot of the HRS web portal

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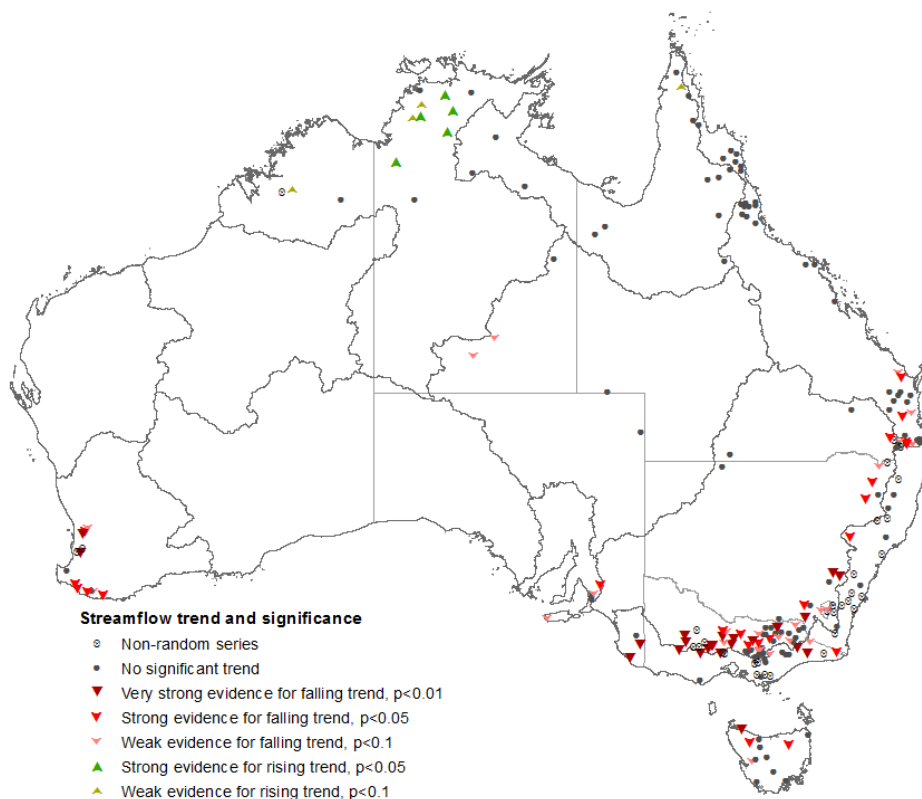


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650 Figure 4: Stacked bar-plot summarizing the MK trend test results for the 222 HRS stations,
 651 with data labels showing the number of stations in each category (Q_T: annual total flow,
 652 Q_BF: annual baseflow, Qmax: daily maximum flow, Q90: 90th percentile daily flow, Q50:
 653 50th percentile daily flow, Q10: 10th percentile daily flow, Q_DJF: summer flow, Q_MAM:
 654 autumn flow, Q_JJA: winter flow, Q_SON: spring flow)

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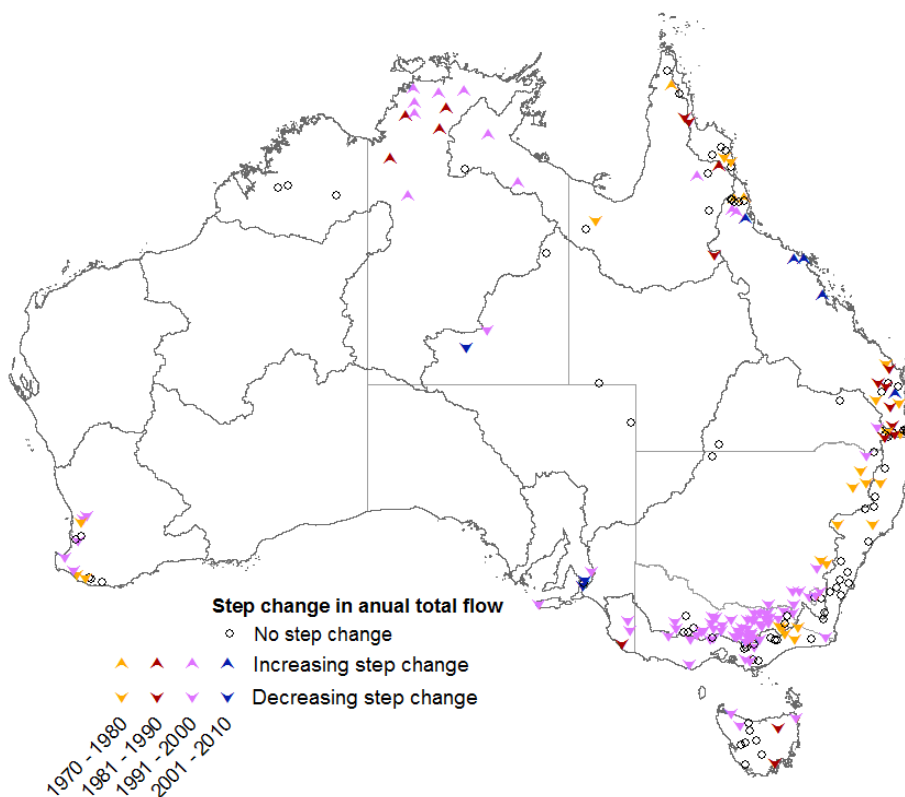


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658 Figure 5: Spatial variation in trend results of annual total flow, decrease trends were shown in
659 significance levels at 0.01, 0.05, and 0.1

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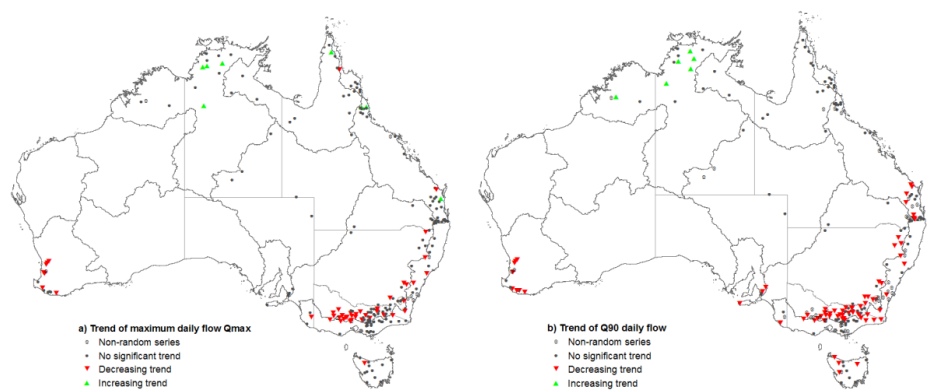
663 Figure 6: Variations of step change in annual total flow for stations showing significant

664 increase or decrease trend

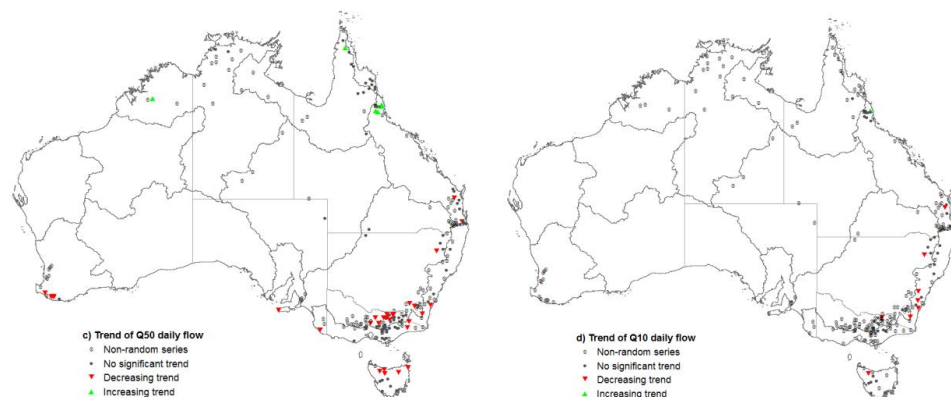
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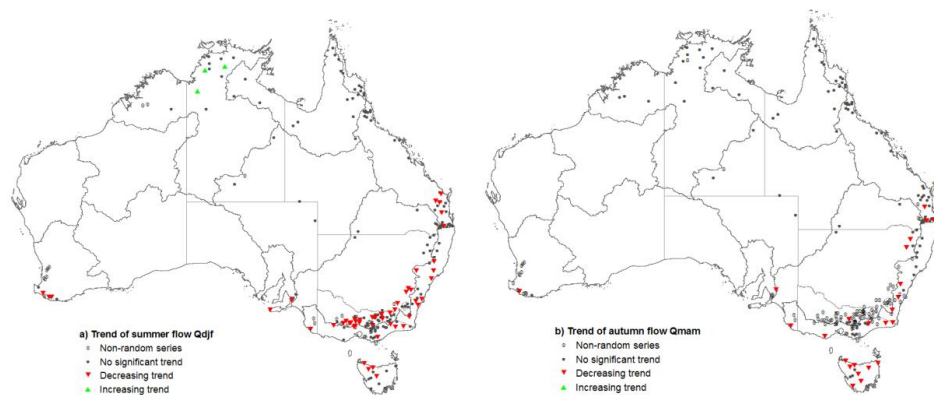
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669 Figure 7: Maps showing trends of daily flow in various magnitude categories a) maximum
670 daily flow Q_{Max} ; b) Q_{90} daily flow; c) Q_{50} daily flow; d) Q_{10} daily flow at 10% significant
671 level ($p < 0.1$)

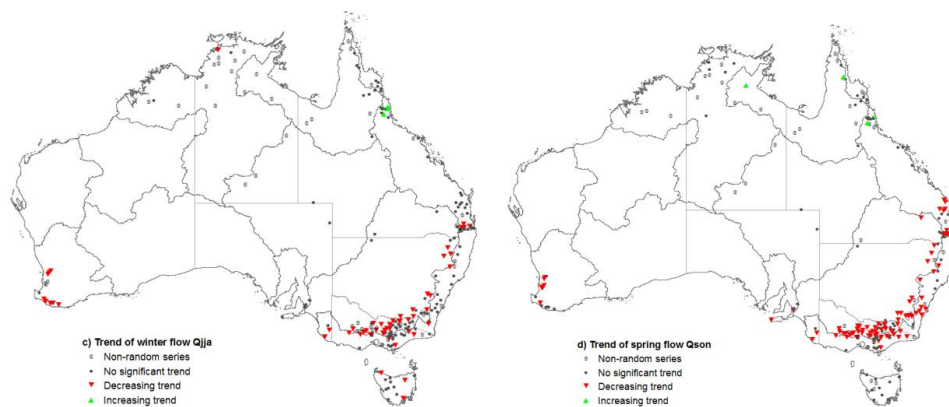
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676 Figure 8: Maps showing trends of seasonal flow in a) Summer; b) Autumn; c) Winter; d)

677 Spring