Real-time monitoring of nitrate transport in deep vadose zone under a crop field—implications for groundwater protection

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Abstract

Nitrate is considered the most common non-point pollutant in groundwater. It is often attributed to agricultural management, when excess application of nitrogen fertilizer leaches below the root zone and is eventually transported as nitrate through the unsaturated zone to the water table. A lag time of years to decades between processes occurring in the root zone and their final imprint on groundwater quality prevents proper decision-making on land use and groundwater-resource management. This study implemented the vadose monitoring system (VMS) under a commercial crop-field. Data obtained by the VMS for of 6 years allowed, for the first time known to us, a unique detailed tracking of water percolation and nitrate migration from the surface through the entire vadose zone to the water table at 18.5 m depth. A nitrate concentration time series, which varied with time and depth, revealed—in real time—a major pulse of nitrate mass propagating down through the vadose zone from the root zone toward the water table. Analysis of stable nitrate isotopes indicated that manure is the prevalent source of nitrate in the deep vadose zone and nitrogen transformation processes have little effect on nitrate isotopic signature. The total nitrogen mass calculations emphasized the nitrate mass migration towards the water table.

Furthermore, the simulated pore-water velocity through analytical solution of the convection–dispersion equation shows that nitrate migration time from land surface to groundwater is relatively rapid, approximately 5.9 years. Ultimately, agriculture land uses, which are constrained to high nitrogen application rates and coarse soil texture, are prone to induce substantial nitrate leaching.
*Keywords:* Nitrate transport, Deep percolation, Vadose zone, Groundwater pollution
1 Introduction

Groundwater contamination by nitrate originating from agricultural land use is a global problem. The World Health Organization guideline for maximum level of nitrate in the drinking water is 50 mg L\(^{-1}\) as NO\(_3\) (WHO, 2011). The US Environmental Protection Agency (EPA) regards nitrate as requiring immediate action whenever its concentration exceeds drinking-water standards (US EPA, 1994). A detailed framework was established by the Nitrate Directive of the EC (European Community, 1991) to prevent water pollution by nitrate. Nevertheless, nitrate contamination has disqualified drinking-water wells in Israel (local standard: 70 mg L\(^{-1}\) NO\(_3\)) more than any other contaminant at the beginning of the 21st century (Elhanany, 2009). To prevent excessive leaching of nitrate and its arrival to the groundwater, it is essential to investigate and quantify the mechanisms controlling nitrate migration in the unsaturated zone with respect to the specific practices used on agricultural land.

Nitrate fate in the subsurface has been investigated by various approaches, such as (i) isotopic signature analysis in groundwater systems (Oren et al, 2004; Wassenaar et al., 2006; Showers et al., 2008; Baram et al., 2013), (ii) crop-management strategies, which combine crop production and nitrate leaching to the subsurface (Hanson et al., 2006; Doltra and Muñoz, 2010; Beggs et al., 2011), and (iii) studies based on data from the deep vadose zone (Dann et al., 2010; Nolan et al., 2010; Botros et al., 2012; Kurtzman et al., 2013; Dahan et al., 2014; Turkeltaub et al., 2015b). Nevertheless, estimates based on data obtained from excavated soil profiles and pore-water sampling during a short period of time represent a snapshot in time of the sediment’s chemical state rather than dynamic temporal variations. Moreover, the drawback of methods based on frequent groundwater sampling from wells is that the
concentration of nitrate might already be at levels that will lead to disqualification of the aquifer as a source for drinking water.

The transfer time of nitrate within the deep vadose zone has been estimated to take from weeks to decades, depending on the water regime, thickness of the unsaturated zone and lithological characteristics of the subsurface (Spalding et al., 2001; Scanlon et al., 2010). Knowledge of nitrate's fate and transport below the root zone is restricted due to issues such as soil spatial variability and long travel times in the deep vadose zone (Onsoy et al., 2005). Moreover, estimates of cumulative nitrate fluxes in the unsaturated zone have shown significant differences in the timing and magnitude of fluxes derived from different land uses (Green et al., 2008; Dahan et al., 2014; Turkeltaub et al., 2014, 2015b). Our understanding of the cumulative effect of nitrate leaching from the root zone through the unsaturated zone on nitrate levels in the groundwater is blurred by mixing and dilution in the aquifer water. The tendency toward elevated nitrate concentration in aquifer water is thus a relatively slow process (Green et al., 2008). Knowing the time lag between initiation of a pollution process in the unsaturated zone and its final effect on aquifer quality could give decision-makers more time to plan possible backups for alternative water supply (Baram et al., 2014).

The recent development of a vadose-zone monitoring system (VMS) enables continuous monitoring of the hydrological and chemical properties of percolating water in the deep vadose zone under agriculture settings (Turkeltaub et al., 2014, 2015b) and other hydrological settings (e.g. Dahan et al., 2009; Baram et al., 2013). Data collected by the system comprise direct measurements of the water-percolation fluxes and the chemical evolution of the percolating water across the entire unsaturated zone. An earlier investigation at the present study site implemented the VMS and demonstrated the percolation patterns, chloride accumulation and
groundwater recharge behavior and tendency in the deep vadose zone of two agricultural settings, a grapefruit orchard and a crop field (Turkeltaub et al., 2014). Unsaturated flow models were calibrated to the water content observation and were used for groundwater recharge fluxes simulations.

The objective of the present study was to demonstrate the water flow and nitrate transport through the deep vadose zone underlie the crop field, with respect to rain patterns as well as the agricultural and fertilization setup. Continuous data on variations in the sediment water content and nitrate concentrations were collected from the entire vadose zone for over 6 years. The nitrate concentration time series, which included variation of nitrate in time and at multiple depths, revealed, in real time, a major pulse of nitrate mass propagating down through the vadose zone toward the water table. These results indicate that nitrate fluxes in the unsaturated zone underlie agriculture land-uses were associated with high nitrogen application rates and coarse texture soils. Furthermore, pollution events originated from agriculture land-uses can be monitored in their early stages, long before pollution accumulates in the aquifer water.

2 Methods

2.1 Study area

A commercial crop field site was selected as a representative prevalent agriculture setting in the southern part of the coastal plain of Israel (34°41’13” E; 31°49’42” N) and is part of an array of VMSs that were installed under different representative land-uses situated above the southern part of the phreatic coastal aquifer.
The study was conducted between 09/2009 and 04/2015. Mediterranean climate prevails in this area, with hot, dry summers (May–September) and rainy winters (October–April), with an average annual rainfall of 512 mm and average temperatures of 31.2 °C (August) and 17.8 °C (January) in the hottest and coldest months, respectively (Israeli Meteorological Service, 2015). Reference evapotranspiration rates calculated according to the Penman–Monteith method (suggested by the Food and Agriculture Organization) range from 1.5 mm day\(^{-1}\) (January) to 5.7 mm day\(^{-1}\) (July) (Israeli Meteorological Service, 2015).

The crop field cultivation history includes alternation between rainfed agriculture, as wheat and irrigated agriculture as watermelon for seeds and cotton as summer crop (personal communication). From 2005 to 2013, the crop field site was cultivated with rainfed winter crops—spring wheat (*Triticum aestivum* L.) and pea (*Pisum sativum* L.) (Fig. 1). Then for 1 year (2013/2014), the field was uncultivated. The crops were sown at the beginning of the wet season (November) and grew into the spring (April). After harvest, disk plow and roller practices were implemented. Since 2005, the main fertilization application to the field was dairy-farm slurry manure, which was distributed over the 10 ha field for 60 days during May and June (Fig. 1). The total nitrogen concentration in the dairy slurry is 900 mg L\(^{-1}\) (Water Authority, 2012). In September 2014, jojoba (*Simmondsia chinensis*) shrubs were planted and irrigation systems were installed.

### 2.2 Monitoring
The field was instrumented with a VMS in May 2008 (Fig. 1). Full technical descriptions of the VMS structure, performance and installation procedures can be found in other publications (Rimon et al., 2007, 2011; Dahan et al., 2008, 2009). For brevity, only a general description is given here.

The VMS is composed of a flexible sleeve installed in an uncased, slanted (35° to the vertical) borehole hosting multiple monitoring units at various depths. Each monitoring unit consisted of a flexible time-domain reflectometry (FTDR) sensor for continuous measurements of sediment water content, and a vadose-zone sampling port for frequent collection of pore-water samples from the unsaturated zone (Table 1). The slanted installation ensures that each monitoring unit faces an undisturbed sediment column that extends from land surface to the probe or sampling port depth. After insertion of the VMS into the borehole, the flexible sleeve was filled with a high-density solidifying material (liquid two-component urethane) that solidifies in the borehole shortly after its application, thereby ensuring proper sleeve expansion for good contact of the monitoring units with the borehole’s irregular walls, sealing its entire void and preventing potential cross-contamination by preferential flow along the borehole.

Since each monitoring unit is located under its own undisturbed sediment column, the integrated data from the VMS should be regarded as representative of a wider zone rather than a single vertical profile. Sediment water content was monitored daily. Pore-water sampling from the unsaturated sediments is achieved by creating hydraulic continuity between the sediment and the sampling port using a flexible porous interface (Dahan et al., 2009; Patent # US 6,956,381; US 12/222,069; EP 07706061.4; IL 193126). The vadose zone sampling ports (VSPs) are operated through a set of small-diameter access tubes and control valves. Prior the water
sampling collection, a low pressure (vacuum) is applied to the sampling ports to draw
the sediment pore water. Subsequently, the water samples are retrieved using
pressurized gas (N\textsubscript{2}) to push the sample to the surface. Water samples were collected
every 90 days on average, from 09/2009 to 04/2015. Samples were stored chilled in
the field and at 4°C in laboratory after filtered through a 45 µm filter. Chemical
analyses were performed the following day. The monitoring system operated with
Campbell Scientific (Logan, UT) data acquisition and logging instruments, including
TDR100, SDM50X, AM 16/32 multiplexers and a CR10X datalogger.

2.3 Chemical and Isotopic Analyses

Nitrate and chloride concentrations in the water samples were determined using ion
chromatography (DIONEX, 4500I). The isotopic composition of nitrate 15N and 18O
in the water samples was determined through nitrate reduction to nitrogen dioxide,
which was then analyzed using a gas mass spectrometer (McIlvin and
Altabet, 2005).

2.4 Nitrate-transport simulations

The observed nitrate concentration dynamics at the 6.3 m, 9.5 m, 15.6 m and
18 m depths (Table 1) were analyzed and compared with earlier modeling estimations
conducted according to observations of water content under the crop field (Turkeltaub
et al., 2014). Nitrate transport was modeled in terms of the convection–dispersion
equation (CDE) equilibrium assuming resident concentration for a third-type inlet
condition as follows (Toride et al., 1999):

\[
R \frac{\partial c}{\partial t} = D \frac{\partial^2 c}{\partial x^2} - v \frac{\partial c}{\partial x} \tag{1}
\]
where \( c \) is the solute concentration, \( x \) is distance, \( t \) is time, \( D \) is the dispersion coefficient, \( v \) is the average pore water velocity (water flux \( q \) divided by the water content \( \theta \)), and \( R \) is the retardation factor.

The nitrate concentrations obtained by the VSP at the 4.2 m depth (Table 1) served as a series of successive applications of solute pulses (multi-pulse boundary condition). All of the sampling ports are located in a relatively homogeneous medium of sandy texture (Turkeltaub et al., 2014), following the intrinsic assumption of CDE analytical model homogeneity. The CXTFIT2 code (Toride et al., 1999) and the Levenberg–Marquardt-type optimization approach (Marquardt, 1963), both included in STANMOD (van Genuchten et al., 2012), were used for inversely estimating the pore-water velocity (\( v \)) and dispersion coefficient (\( D \)) according to observed concentrations. Both parameters were obtained by running CXTFIT2 multiple times for inverse optimization, each time with different initial values (Turkeltaub et al., 2015a,b).

### 2.5 Total nitrate mass

The total nitrate mass in the unsaturated zone estimations was calculated to emphasize the nitrate mass that will eventually contaminate the groundwater. The following equation was used for yearly nitrate mass (per area) in the vadose zone:

\[
M = \int_{Z=\text{ground surface}}^{Z=\text{water table}} \bar{\theta}_i \times C_i \times dz_i
\]

where \( M \) is nitrate mass in the vadose zone under a unit area, \( i \) indexes the depth interval for which the corresponding sampling port is at its centre, \( C_i \) is the nitrate concentration [M L^{-3}] sampled with the sampling port at that depth interval, \( \bar{\theta}_i \) is the
average water content measured by the nearest FTDR sensor \([L^3 L^{-3}]\), and \(dz_i\) is the interval length [L] (Fig. 2).

3 Results and discussion

3.1 Nitrate migration in the unsaturated zone

The continuous monitoring of the vadose zone show temporal variations in measured water content (Fig. 2). Throughout the monitoring period, most of the rainstorms caused a rise in the water content measured by the shallowest water sensor (0.5 m, Fig. 2). At the 2.1 m and 3.1 m depths, the rise in water contents corresponded mainly to larger rain events (Fig. 2b,c). The sensors at the deeper depths displayed temporal variability with respect to the cumulative annual rain pattern. In some years, a lag between the end of the rainy season and the rise in water content was recorded, whereas in other years, the rise in water content occurred throughout the entire vadose zone following a significant rain event (Fig. 2d–h). A more detailed description of the sequential rise in water content with depth following a wetting event on land surface, and a clear indication of propagation of a wetting wave through the vadose zone are presented in our earlier study at the site (Turkeltaub et al., 2014), and in other studies at different sites (Rimon et al., 2007, 2011; Dahan et al., 2008, 2009; Baram et al., 2012, 2013).

Throughout 6 years of continuous monitoring, variations in nitrate concentration were observed (Fig. 3). The nitrate concentration time series with depth (Fig. 3) reveals a major pulse of elevated concentrations, initiating close to the surface in 2011 and 2012, and gradually progressing down the vadose zone toward the water
table at a depth of about 18 m. The process was first monitored at the uppermost sampling port at 1 m depth, where nitrate concentrations displayed a significant increase during the winter of 2010/2011. Then a gradual trend of reduction in nitrate concentration was observed at this depth until March 2014. A close examination of the nitrate concentrations at 1 m depth indicated repeating fluctuations, with higher nitrate concentrations after harvest due to application of the dairy slurry, and then followed by a reduction in concentrations. Although hard to notice at the illustrated scale in Fig. 3a, the nitrate concentrations between September 2009 and September 2010 were still relatively high and fluctuated near 600 mg L$^{-1}$ (Fig. 3a). Then they escalated to about 3200 mg L$^{-1}$ after cultivation of the pea crop. Following this relatively large increase in nitrate concentration in May 2011, a decline was observed until January 2012 to about 1500 mg L$^{-1}$ (Fig. 3a). This phenomenon repeated itself in April 2012, when the nitrate concentration increased again to 2800 mg L$^{-1}$ and then decreased to 78 mg L$^{-1}$ in April 2015 due to cessation of slurry application (Fig. 3a, note the solid line arrow).

The distributed estimated nitrogen mass over the field is approximately 200 Kg ha$^{-1}$ year$^{-1}$, which is in the range of the European application recommendations (van Grinsven et al., 2012). The Agriculture Extension Service of Israel (2016) recommendation concerning nitrogen fertilizer application for wheat crop (main crop) is between 40 and 100 kg ha$^{-1}$. Therefore, an excessive amount of nitrogen is applied by disposing dairy wastes over the field. Moreover, nitrogen fixing agents in agricultural systems are the symbiotic associations between legumes and rhizobia (Rochester et al., 2001). Rotation between legume crop and non-legume crop practice supposes to replace some of the need in nitrogen fertilizer (Rochester et al., 2001). The average nitrogen fixation by pea crop, according to global data sets, is 86 Kg ha$^{-1}$
year\(^1\) (Herridge et al., 2008), which is about 43% of the nitrogen applied by the dairy slurry. Thus, application of dairy farm slurry combined with a legume crop (pea) seemed to have enriched the top soil with excess nitrogen, as compared to cultivation of cereal-type crops (Fig. 3a).

Progression of the nitrate migration deeper into the vadose zone can be divided into two periods. In the first period, October 2010 to January 2013, at depths of 2.7, 4.2, 9.5 and 15.6 m (Fig. 3b,c,e,g), the increase in nitrate concentration was moderate and continuous\(^1\), whereas at depths of 6.3 and 18 m, there was no major change in nitrate concentrations (Fig. 3b-d). In the second period\(^1\) starting from July 2013 following the rainy winter of 2012/13, substantial nitrate breakthroughs were noticeable throughout most of the vadose zone cross section (marked with arrows in Fig. 3). This rapid nitrate progression to the deeper parts of the vadose zone could be related to the soil's physical characteristics. In the top 3 m, the soil comprised of fine-textured layers (sandy-loam and loamy sand), and from 3 to 18.5 m (water table), the soil consisted of a coarser sand-textured layer (Turkeltaub et al., 2014). Thus, as a consequence of substantial water percolation, which induced intensive water flux across the coarse-textured soil, nitrate transport could be detected at deeper depths of the vadose zone.

Here, as well in previous studies in literature, nitrate fluxes in the unsaturated zone underlie agriculture land-uses were associated with nitrogen application rates and soil physical properties (Green et al., 2008; Botros et al., 2012; Turkeltaub et al., 2015b). Therefore, to attenuate nitrate leaching to aquifers, search should be dedicated to locate the ‘hot spots’ where these conditions prevailed (Liao et al., 2012).

3.2 Nitrate sources
The $\delta^{15}$N values clearly showed that manure is the main source of nitrate in the vadose zone pore water (Fig. 4). Nitrate isotope composition in the vadose zone pore water depends on nitrogen sources and transformation processes (Böhlke, 2002). Examination of the isotopes values suggested that transformation processes such as denitrification and mineralization of soil nitrogen sources have little effect on nitrate isotopic signature. As discussed in the previous section, the relatively rapid nitrate transport downward to deeper parts of the vadose zone is controlled by soil properties and nitrogen application rates. These factors reduce the potential for transformation processes and plant uptake to occur (Liao et al., 2012). Moreover, Various studies conducted under similar conditions (soil types and agriculture land use) as in the current study, presented insignificant nitrogen transformation processes and doubt the ability of attenuating nitrate within the deeper vadose zone (Green et al., 2008; Burow et al., 2010; Gautam and Iqbal 2010; Dann et al., 2013; Zhang et al., 2014; Turkeltaub et al., 2015b). Yet, other studies suggested contrast conclusions. Salazar et al. (2012) reported on low nitrate leaching rates in spite of high nitrogen application rates and Lockhart et al. (2013) claimed that depth to groundwater provided a significant control on nitrate concentration in groundwater regardless of soil type or crop type. Thus, a holistic approach comprises all potential factors that control nitrate fluxes to groundwater should be held to identify the dominant ones.

3.3 Nitrate storage in the vadose zone

The yearly nitrate mass calculations (Eq. 2) displayed an increase from 2009 to 2010 (Fig. 5), at the same time as NO$_3^-$ concentration increased in the upper part of the vadose zone (Fig. 3a). Subsequently, the highest increase in nitrate mass was calculated for 2011 following the combination of cultivation of the pea crop and
excessive application of dairy slurry (Fig. 5). It seems that the yearly fluctuations in
calculated nitrate mass can be explained by the lag time in the transport process
between the sampling points. Hence, the peak in nitrate mass observed in the upper
parts during 2011 remained in the vadose cross section and eventually reached the
deeper parts of the vadose zone as a breakthrough type (Fig. 5).

3.4 Nitrate transport model

Using nitrate time series obtained from deeper part of the vadose zone for
model simulations allowed avoiding the highly dynamic nature of the root zone.
Furthermore, transport calculations are less effected by mass balance uncertainties as
according to previous section, nitrate attenuation processes are insignificant in deep
vadose zone.

The results indicated relatively good agreement between observed and
simulated nitrate concentration trends (Fig. 6). Nevertheless there were discrepancies
in the absolute values and with the simulated nitrate concentrations increasing before
the observed concentrations at the 6.3 and 18 m depths (Fig. 6a, d). These gaps could
be explained by the assumptions that are intrinsic to the CDE model (Eq. 1) —
homogeneous medium and average velocity—along with the assumption of even
distribution of the nitrogen source on the surface. Nevertheless, the CDE provided an
approximation that could be compared with earlier numerical modeling results (van
Genuchten et al., 2012). The calculated hydrodynamic dispersion coefficient was 81
cm$^2$ day$^{-1}$ and the pore-water velocity was 0.836 cm day$^{-1}$, which is about 305 cm
year$^{-1}$. Multiplying the velocity by the weighted average water content, 0.060 cm$^3$ cm$^{-3}$
(Fig. 2c-h), the Darcian flux equaled 18.3 cm year$^{-1}$, which is very similar to earlier
average flux estimation of 19.9 cm year$^{-1}$ averaged for 24 years (Turkeltaub et al.,
If neglecting the diffusion term in the hydrodynamic dispersion coefficient, the estimated longitudinal dispersivity ($D/v$) is 97 cm. The calculated dispersivity value is relatively large compared with reported values from earlier solute transport investigations in sandy texture soils (e.g. Toride et al., 2003; Dann et al., 2010). However, it was showed that dispersivity increases with travel distance (Vanderborght and Vereecken, 2007).

The calculated nitrate transport time from land surface to groundwater is approximately 5.9 years. Yet, the increase in nitrate concentration at the 18 m depth occurred in July 2013, which is 8 years after the first slurry application. Olson et al. (2009) reported that there was a threshold amount of slurry application before nitrate accumulated in the soil. Hence, the gap of 2 years between the first application and nitrate arrival to 18 m depth might be related to the period before critical amount of manure was applied to the field.

3.5 Practical implications of vadose-zone monitoring

To prevent a long-term gradual degradation in groundwater quality, the link between sources of pollution on the surface and their migration pattern in the unsaturated zone should be understood long before their final cumulative imprint in the aquifer water. Herein, the application of a VMS under an agricultural field enabled, for the first time known to us, real-time tracking of water flow and nitrate transport from the surface through the entire deep vadose zone. Accordingly similar monitoring concepts for the vadose zone can be used as an alert apparatus for pollution events in their early stages while pollution is still migrating in the unsaturated zone, and long before accumulation in the aquifers water.
This study demonstrates how nitrate concentrations in the vadose zone exceed the local standard for disqualified drinking-water wells and threaten the groundwater quality. Hence, agro-hydrologically sustainable manure application rates, i.e. sufficient crop production and minimizing nitrate leaching, could be satisfied by suitable regulation or adjustments to meet crop requirements (Olson et al. 2010). To optimize the efficiency of the manure distribution methodology, estimations should include the controlling factors as soil properties, crop type, season, nitrogen attenuation processes and the critical amount of manure application before nitrate accumulation in the soil occurs. Considering only part of the factors could lead to the opposite result. For example, the manure application in this study occurred during the beginning of the dry period, May and June (there are no rain events till October) to prevent nitrogen leaching due to rain events. However, the distributed nitrogen was retained in the soil till winter time and did not undergo significant attenuation processes. The incorrect assumption of manure distribution during the dry period resulted in intensive nitrate leaching. Furthermore, according to the observations presented in this study, the manure application should be reduced following legume crop type. Yet, in many cases, there is a surplus amount of manure to be disposed. Therefore, alternative methods for waste management have to be utilized, coincided with regulating manure application (Westerman and Bicudo, 2005; van Grinsven et al., 2012).

Nitrate transport from land surface to water table through a relatively thick vadose zone occurred within less than a decade. This is a considerably rapid pollutant migration when considering remediation strategies. Moreover, the nitrate observations obtained by the VMS and the isotopic signature analysis indicated that nitrate attenuation processes are insignificant. Hence, agriculture sites constrained to similar
conditions as in this study, most of the nitrate mass that leaches under the root zone will eventually reach groundwater.

4 Summary and Conclusions

An intensive nitrate leaching beyond root zone was attributed to soil properties and nitrogen application rates. The implementation of a vadose zone monitoring system (VMS) under an agricultural field enabled real-time tracking of water flow and migration of a nitrate plume from the surface through the deep vadose zone to the water table at 18.5 m depth. Isotopic composition of nitrate-nitrogen in the water samples indicated that manure is the main nitrogen source for nitrate in the vadose-zone pore water. Nitrogen transformation processes seem to have only little effect under an intensively fertilized crop field. Total nitrate mass estimations displayed the nitrate mass advancement toward the deep vadose zone. Moreover, according to the simulated pore-water velocity, nitrate arrival to water table occurred within less than a decade.

As in this study, an array of VMSs was installed under other representative agriculture land-uses situated above the southern part of the Israeli coastal aquifer. The findings from each site are combined to generate a comprehensive perspective on dominant factors controlling groundwater quality and quantities. Subsequently, these conclusions will be examined with a regional scale aquifer transport model.

Protection of groundwater from potential pollution originating from agricultural land uses has to include effective and continuous monitoring of the vadose zone. Pollution events can be monitored in their early stages, long before pollution accumulates in the aquifer water.
Acknowledgements

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References


(Available at


Table 1

Depth distribution of the vadose-zone monitoring system (VMS) units.

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<th>Vertical depth from land surface (m)</th>
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1 Flexible time-domain reflectometry probe.
Figures

Figure 1. Crop field site with monitoring location during two periods: crop growth during the wet season (a), and after harvesting and during slurry application (b). (c) Schematic illustration of the vadose-zone monitoring system installed under the crop field. *Vadose zone sampling port*, vadose-zone sampling port; FTDR, flexible time-domain reflectometry sensor.

Figure 2. Water-content ($\theta$) at different depths in the vadose zone and daily rainfall for six consecutive years.

Figure 3. Time series of observed ($NO_3$) concentrations in the vadose zone and daily rainfall for six consecutive years.

Figure 4. $\delta^{15}N$ profile of nitrate in the water samples obtained from the vadose zone under the crop field.

Figure 5. Yearly total nitrate mass of the entire vadose zone per year of sampling.

Figure 6. Observed (red dots) and simulated (dash blue line) nitrate concentrations for the vadose-zone sampling port at the 6.3 m, 9.5 m, 15.6 m and 18 m depths. Nitrate concentration series from each depth served as a multiple pulse input boundary condition to the consecutive depth.
Vadose zone

Groundwater

18.5 m